

## POPULAR SUMMARY FOR THE ARTICLE

“Saharan Air and Atlantic tropical cyclone suppression from a global modeling perspective.”

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being submitted to *Geophysical Research Letters*

During summer 2006, the NASA African Monsoon Multidisciplinary Analysis (NAMMA) organized a field campaign in Africa called Special Observation Period (SOP-3), in which scientists in the field were involved in a number of surface network and aircraft measurements. One of the scientific goals of the campaign was to understand the nature and causes for tropical cyclogenesis originating out of African Easterly Waves (AEWs, westward propagating atmospheric disturbances sometimes associated with precursors of hurricanes), and the role that the Saharan Air Layer (SAL, a hot and dry air layer advecting large amounts of dust) can play in the formation or suppression of tropical cyclones. During the NAMMA campaign a high-resolution global model, the NASA GEOS-5, was operationally run by the NASA Global Modeling and Assimilation Office (GMAO) in support to the mission. The daily GEOS-5 forecasts were found to be very useful by decision-making scientists in the field as an aid to discriminate between developing and non-developing AEWs and plan the flight tracks.

In the post-event analyses which were performed mostly by the Goddard Laboratory for Atmospheres, two events were highlighted: a non-developing AEW which appeared to have been suppressed by Saharan air, compared to a developing AEW which was the precursor of hurricane Helene. Both events were successfully predicted by the GEOS-5 during the real-time forecasts provided in support to the mission.

In this work it is found that very steep moisture gradients and a strong thermal dipole, with relatively warm air in the mid-troposphere and cool air below, are associated with SAL in both the GEOS-5 forecasts and the NCEP analyses, even at -great distance- from the Sahara.

The presence of these unusual thermodynamic features over the Atlantic Ocean, at several thousands of kilometers from the African coastline, is suggestive that SAL mixing is very minimal and that the model's capability of retaining the different properties of air masses during transport are important to represent effectively the role of dry air intrusions in the tropical circulation.

# Saharan Air and Atlantic tropical cyclone suppression from a global modeling perspective.

O. Reale,<sup>1,2</sup> W. K. M. Lau,<sup>1</sup> A. daSilva,<sup>3</sup> K.-M. Kim<sup>1,2</sup>

This article investigates the role of the Saharan Air Layer (SAL) in two cases of non-developing and developing systems observed during the Special Observation Period (SOP-3) phase of the 2006 NASA African Monsoon Multidisciplinary Analyses (NAMMA). A high-resolution global model, the NASA GEOS-5 was operationally run by the NASA Global Modeling and Assimilation Office (GMAO) in support to the NAMMA field campaign, which included surface network and aircraft measurements. One of the scientific goals of the campaign was to understand the nature and causes for tropical cyclogenesis out of African Easterly Waves (AEWs). The daily GEOS-5 forecasts were found to be very useful by decision-making scientists in the field as an aid to discriminate between developing and non-developing AEWs and plan the flight tracks. In the post-event analyses of a non-developing system which appeared to have been suppressed by SAL, it has been found that very steep moisture gradients and a strong thermal dipole are associated with SAL in both the GEOS-5 forecasts and the NCEP analyses, even at great distance from the Sahara. This is suggestive that SAL mixing is very minimal and that the model's capability of retaining the different properties of air masses during transport is important to represent effectively the role of dry air intrusions in the tropical circulation.

## 1. Introduction

African Easterly Waves (AEWs) have been recognized as prominent weather-producing events of northern tropical Africa [e.g Burpee, 1974; Asnani, 2005] and have been extensively studied from observational and modeling perspectives [e.g Hsieh and Cook, 2005; Kiladis et al., 2006].

However, the development of AEWs into tropical depressions remains one of the most challenging problem in the prediction and modeling of Atlantic tropical cyclones. The Saharan Air Layer (SAL), a layer of hot dry air rich in dust and produced over the Saharan desert, has been investigated with the aid of Geostationary Operational Environmental Satellite (GOES) by Dunion and Velden [2004] and recognized as a possible, important mechanism

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43 concurring to tropical cyclone suppression.

44 The general problem of tropical cyclogenesis has al-  
45 ways been considered from either strictly observational  
46 or high-resolution mesoscale points of view, because the  
47 lower resolution of global models is deemed to be inad-  
48 equate to trigger spontaneous cyclogenesis. However, in  
49 recent years a number of global models have reached the  
50 resolution of 10-40 km and have started to display some  
51 tropical cyclogenesis capability [e.g Atlas et al., 2005;  
52 Shen et al., 2006].

53 The organization of several convective centers into  
54 a rotating system still requires cloud-resolving models  
55 and can not be reproduced by real-time global numer-  
56 ical weather prediction models. However, the process of  
57 cyclogenesis or cyclogenesis suppression can be studied  
58 from different perspectives, and global models do have  
59 the advantage of better capturing the large-scale forcings  
60 involved. In particular, a global model can be used to  
61 investigate the role of SAL not only on the wave scale,  
62 but also from the point of view of the large-scale trans-  
63 port from its source region, and can therefore represent  
64 the possible modification of thermodynamical properties,  
65 as the waves propagate over thousands of kilometers. At  
66 the same time, sufficiently high resolution is needed to  
67 unveil some of the SAL kinematic features, such as the  
68 increasingly narrow structure of the dry air filaments be-  
69 ing intruded in a tropical circulation, and the sharpness  
70 of the boundaries between Saharan and non-Saharan air.

## 2. The Model

71 In this work, the effect of Sahara air is investigated  
72 using a high-resolution global atmospheric model, the  
73 NASA GEOS-5, documented in Bosilovich et al. [2006].  
74 The GEOS-5 shares the same dynamical core [Lin, 2004]  
75 with the so-called NASA finite-volume General Circula-  
76 tion Model (fvGCM) which has demonstrated remarkable  
77 capabilities in hurricane forecasting [Atlas et al., 2005;  
78 Shen et al., 2006]. The GEOS-5 however contains a new  
79 physics developed predominantly by the Global Model-  
80 ing and Assimilation Office (GMAO), which is substan-  
81 tially different from the previous fvGCM. The version  
82 used during NAMMA was run at a horizontal resolution  
83 of  $0.25^\circ \times 0.33^\circ$  with 72 vertical levels.

## 3. MAP06 and NAMMA

84 Over the past ten years, the NASA Global Modeling  
85 and Assimilation Office, continuing the work previously  
86 done in the Data Assimilation Office, has been increasing  
87 the resolution of its global model and performing real-  
88 time forecasts. During the Atlantic tropical seasons of  
89 2005, 2006 and 2007 the current high resolution version of  
90 the model was put to a severe test by providing real-time  
91 forecasts that could be compared to operational state-  
92 of-the-art models. GEOS-5 was made available to the  
93 operational team in Africa, as an auxiliary decision sup-  
94 port tool in the SOP-3 phase of the NAMMA campaign.  
95 The model performed well and the forecasts were found  
96 beneficial by the team on the field. After-event analysis  
97 has revealed some interesting aspects of the cyclogenetic  
98 process as perceived by the model, and on the impact of  
99 SAL in the development or suppression of storms.

## 4. Analysis

100 In this work we focus on two interesting events ob-

101 served during the SOP-3 NAMMA campaign: one non-  
102 developing and one developing wave, appearing quite sim-  
103 ilar in terms of intensity, vertical shear and other dy-  
104 namical forcings, and we investigate their different evo-  
105 lution and the different properties of the corresponding  
106 SAL intrusions. Figure 1 shows a Hovmöller of 850hPa  
107 relative vorticity and total precipitable water from the  
108 operational NCEP analyses to emphasize AEWs during  
109 the first 15 days of the SOP-3, covering the second half  
110 of August. The strongest wave of the period appears in  
111 the diagram on 23 August at about  $5^{\circ}W$  and undergoes  
112 transition on the following day (hereafter W1). An evi-  
113 dent sharp strip of dry air with the same propagation  
114 speed and amplitude of the wave, clearly associated with  
115 a Saharan Air outbreak, can be seen.

116 In Figure 2, 700 hPa specific humidity and 850 hPa  
117 flow are shown together to emphasize the interaction be-  
118 tween two different levels at the initial time (00z 26 Au-  
119 gust, from NCEP analyses) and across 3 times of the  
120 forecast (24, 48 and 72 hour forecast, corresponding to  
121 verification times of 00z 27, 28 and 29 August respec-  
122 tively). The 700hPa level is at the lower part of the  
123 SAL, whereas 850 corresponds approximately to the top  
124 of the moist lower level and emphasizes the low-level cir-  
125 culation. Flows at 700 and 850 hPa are substantially  
126 different except around the storm center, where a verti-  
127 cally aligned circulation is present from the surface up  
128 to almost 500 hPa. Based upon these forecasts, it was  
129 correctly suggested to the NAMMA forecasting team on  
130 the field that W1 would become a nondeveloping wave  
131 in spite of its apparent strength. The team obviously  
132 had many other forecasting tools available but the infor-  
133 mation provided by the GEOS-5 was correct. Following  
134 the path of the dry air at 700 hPa it can be seen that  
135 as soon as dry air is advected on the top of the 850 hPa  
136 circulation center, the rotating system becomes first elon-  
137 gated and then rapidly evolves into an open wave. This  
138 is even more evident while analyzing intermediate time-  
139 steps (not shown).

140 In Figure 3, a zonal vertical cross-section of specific  
141 humidity at  $20^{\circ}N$  is extracted from the GEOS-5 24-hour  
142 forecast, right across the center of the same circulation  
143 which can still be seen in Figure 2 at 00z 27 August. An  
144 intriguing feature, namely a sharply defined ‘corridor’ of  
145 extremely dry air can be seen down to 800 hPa. Re-  
146 markable moisture gradients are present on both sides.  
147 From the temperature anomaly (obtained by subtracting  
148 the zonal mean between  $80^{\circ}W$  and  $0^{\circ}$ ), a very well-  
149 defined thermal dipole, stronger than any other anomaly  
150 in the range of longitudes selected, can be seen in per-  
151 fect correspondence to the dry tongue. In particular, a  
152 warm anomaly spans between 800hPa and 400hPa, and a  
153 cool anomaly between 825 hPa and the surface. In other  
154 words, since the cross-section cuts across the SAL intru-  
155 sion, it appears that temperature, in the core of the SAL,  
156 is approximately  $3^{\circ}C$  warmer than the surroundings at  
157 the same latitude, in partial agreement with Dunion and  
158 Velden, [2004]. The new aspect of this analysis is however  
159 that a *negative* value also detected in the moist low-level  
160 layer at the base of the column.

161 In Figure 4, the same figure is extracted from the full-  
162 resolution NCEP operational analyses in model levels.  
163 0.5 corresponds roughly to 500 hPa. The Figure confirms  
164 the thermal structure depicted in the GEOS-5 24-hour  
165 forecast. Despite small scale differences, the NCEP anal-  
166 yses confirm the presence of a very well-defined dry intru-  
167 sion at about  $35^{\circ} - 40^{\circ}W$ . Most remarkable is the pres-  
168 ence of the same dipole thermal anomaly seen in Figure 3.

169 A positive value reaching  $4^{\circ}\text{C}$  in the mid-troposphere,  
170 and a corresponding cool anomaly down to  $-4^{\circ}\text{C}$  in the  
171 low moist layer. The anomalies are obtained, as in Fig-  
172 ure 3, by simply subtracting the  $80^{\circ}\text{W} - 0^{\circ}$  mean.

173 In contrast to Fig 1, the Hovmöller computed for  
174 September (Figure s1) shows a powerful wave towards  
175 the end of SOP-3 (hereafter W2) associated to some Sa-  
176 haran Air: however, the magnitude of the SAL is not  
177 comparable with the case of W1.

178 In Figure 5, the same zonal cross-section is produced  
179 across the center of the system W2, which is a precu-  
180 sor of Helene: the vertically aligned vorticity column  
181 at about  $33^{\circ}\text{W}$  is the analyzed signature of the Trop-  
182 ical Storm, named at 00z 14 September 2006 [Brown,  
183 2006]. A weak positive temperature anomaly at about  
184  $40-45^{\circ}\text{W}$  in the lower midtroposphere is associated with  
185 the same Saharan air outbreak which can be detected in  
186 the Hovmöller in Figure S1. However, two prominent dif-  
187 ferences can be seen with respect to the non-developing  
188 W1 in Figures 3 and 4: there is no cool anomaly in the  
189 lowest levels, and the dry air appears more diluted with  
190 less sharp horizontal gradients.

191 The real-time GEOS-5 tropical cyclogenesis forecast  
192 for Helene was correct. In Figure S2, the 850 hPa circula-  
193 tion shows a clearly defined vortex (at about  $24-25^{\circ}\text{W}$   
194 and  $10-12^{\circ}\text{N}$  in the initial conditions) progressing west-  
195 ward and then recurving northwestward, being entangled  
196 in corresponding high levels of 700 hPa moisture. Based  
197 upon this and other information, subjective forecasts on  
198 the field considered the possibility of that system to be  
199 a developing one, and one flight was successfully planned  
200 across it.

201 Post-event model analysis suggests that since vertical  
202 shear and all other environmental conditions were very  
203 favorable in both W1 and W2 case (not shown), but only  
204 the latter underwent development becoming Helene, the  
205 only difference appears to be the intensity of the Saha-  
206 ran Air intrusion. In the model, the temperature dipole  
207 associated to the SAL could be followed at each timestep  
208 and can be considered a possible cause of suppression. In  
209 the precursor of Helene, the low-level negative anomaly  
210 was minimal or absent.

211 While the positive anomaly can be simply attributed  
212 to the signature of warm air originated over the Sahara,  
213 the cool anomaly in the lower levels does not have any  
214 plausible explanation relying on transport only. There  
215 is no source of localized cooler temperatures at that lat-  
216 itude, away from landmass and in a very homogeneous  
217 marine tropical environment. At this time, albeit spec-  
218 ulative, a possible explanation is that the low-level cool  
219 temperatures are an indirect evidence of dust amount.  
220 The thermal effect of Saharan mineral dust is a net reduc-  
221 tion of downwave shortwave radiation in the near-surface  
222 levels, and a heating in the lower midtroposphere, cor-  
223 responding to the core of the SAL. It appears that the  
224 high-resolution NCEP, unlike lower resolution analyses,  
225 can represent this thermal structure and that the GEOS-  
226 5 model initialized by the NCEP analyses could retain  
227 it for 24-72 hours advecting it into the circulation and  
228 producing a realistic cyclone dissipation.

229 In the case of the GEOS-5 forecast, the nature of the  
230 finite-volume dynamics [Lin, 2004] is such that is a par-  
231 ticularly suitable tool to generally maintain sharp gradi-  
232 ents by minimizing unrealistic diffusion processes. The  
233 finite-volume dynamics has been shown to be very ef-  
234 ficient in the midlatitudes where localized temperature  
235 gradients associated with sharp fronts can be very realis-  
236 tically simulated and maintained. This work documents

237 that the same skill can be very useful also in the tropics  
238 when dealing with Saharan Air.

## 5. Concluding Remarks

239 This work documents the contribution of the NASA  
240 Global Modeling and Assimilation Office in support to  
241 the SOP-3 phase of the NAMMA campaign. From the  
242 30 5-day forecasts, one prominent case is extracted, a  
243 very strong non-developing wave, and is compared with  
244 the wave that becomes the precursor of Hurricane He-  
245 lene. GEOS-5 forecasts and NCEP full-resolution analy-  
246 ses document the presence of a strong temperature dipole  
247 (cooler than the environment below 800hPa, and warmer  
248 from 800 to 500hPa) associated with the Saharan air in-  
249 trusion. This dipole is advected into the circulation of the  
250 wave, suppressing further development. No such dipole is  
251 found for the Saharan air intruded in the Helene's precur-  
252 sor. The lower tropospheric cooling associated with the  
253 strong Saharan air outbreak is suggestive that the high  
254 resolution global models and analyses can capture part  
255 of the thermal effect consequent to downward shortwave  
256 reduction caused by large amounts of Saharan dust. An  
257 interactive dust aerosol component, including its direct  
258 radiative effects, is being employed within GEOS-5 to  
259 further assess and quantify the thermal effects of dust on  
260 tropical cyclogenesis and will be used for a future study.

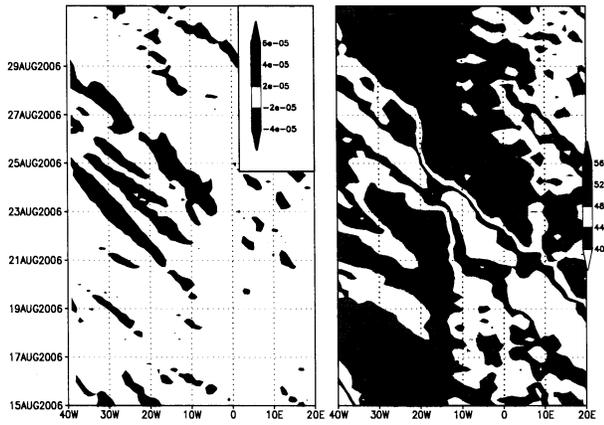
261 **Acknowledgments.** Authors acknowledge support from  
262 Ramesh Kakar through the NAMMA Project. Thanks are also  
263 due to Michele Rienecker for access to the GEOS-5 forecasts.

## References

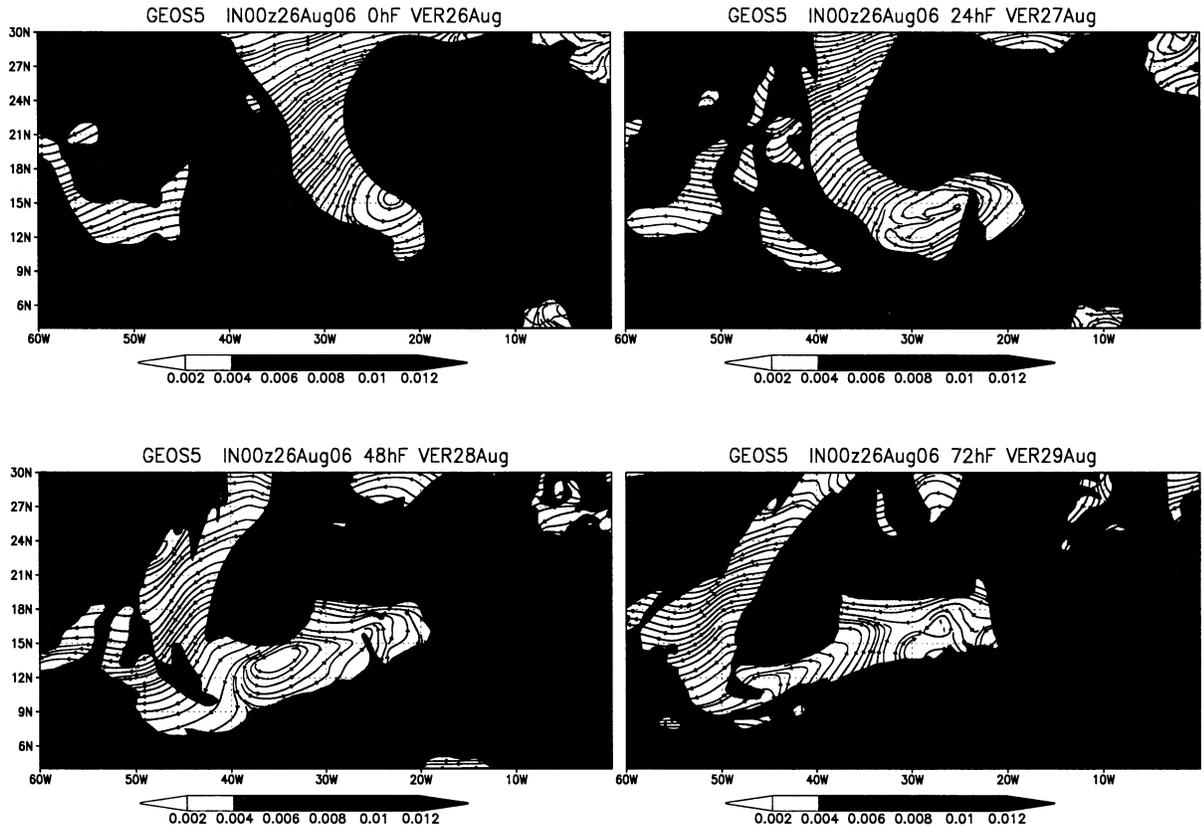
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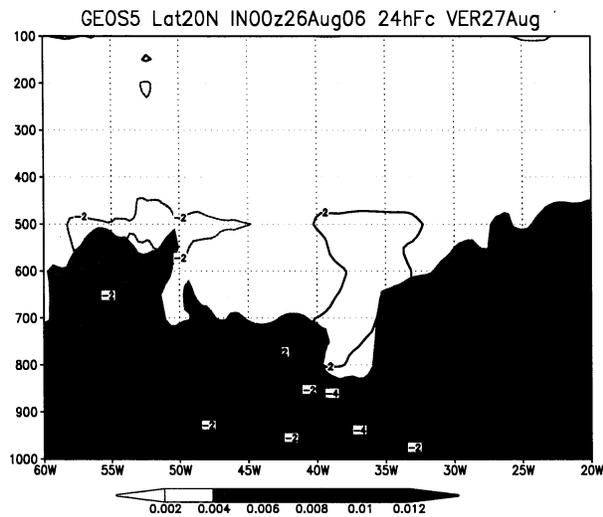
303 O. Reale, Laboratory for Atmospheres, Code 613, NASA  
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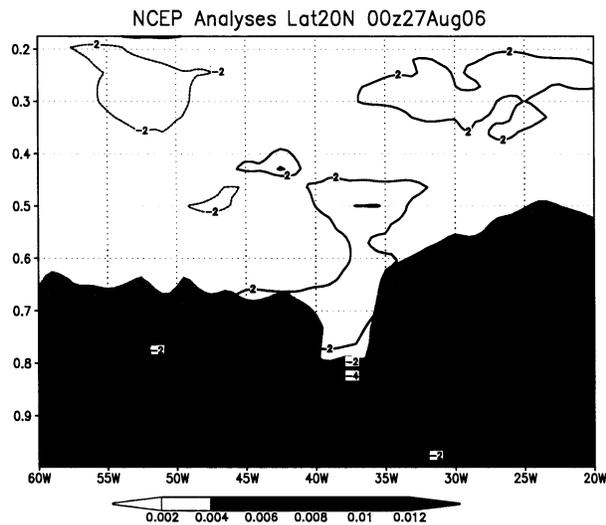
**Figure 1.** Hovmöller of 850 hPa relative vorticity ( $s^{-1}$ , left panel) and total precipitable water (right panel) from the NCEP operational analyses, latitudinally averaged ( $10^{\circ} - 18^{\circ}N$ ), covering the period from 16 to 31 August. Data on pressure levels interpolated on a  $1^{\circ}$  resolution grid. No significant difference from the original data in sigma levels (not shown).



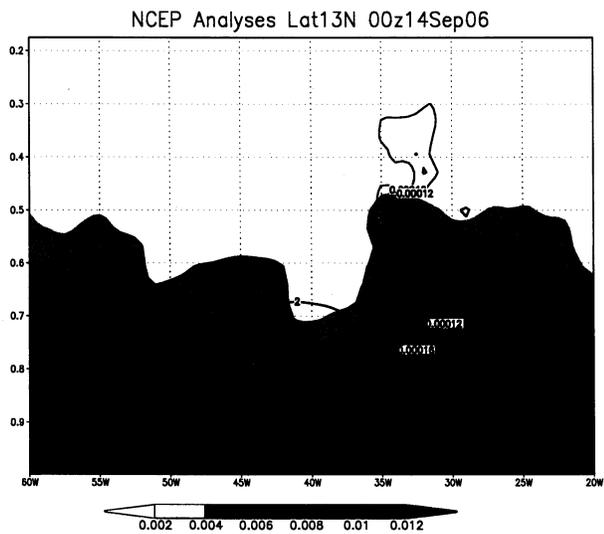
**Figure 2.** GEOS-5: 700 hPa specific humidity ( $KgKg^{-1}$ ) and 850 hPa wind (streamlines) in the NCEP-derived initial conditions (upper left) for 00z 26 August, and relative to the 24, 48 and 72 hour forecasts for 00z 26, 27 and 28 August.



**Figure 3.** GEOS-5: zonal vertical cross-section of specific humidity ( $KgKg^{-1}$ , shaded) and temperature anomaly ( $^{\circ}C$ , contour, subtracting the zonal mean between  $80^{\circ}W$  and  $0^{\circ}$ ) at  $20^{\circ}N$  for 27 August, 24 hour forecast initialized at 00z 26 August.



**Figure 4.** NCEP full-resolution analyses in model levels: zonal vertical cross-section of specific humidity ( $KgKg^{-1}$ , shaded) and temperature anomaly ( $^{\circ}C$ , contour, subtracting the zonal mean between  $80^{\circ}W$  and  $0^{\circ}$ ) at  $20^{\circ}N$  for 00z 27 August. The vertical dimension is only approximately comparable with Figure 3 since the spacing between model levels and pressure levels is different.



**Figure 5.** NCEP full-resolution analyses in model levels: zonal vertical cross-section of specific humidity ( $KgKg^{-1}$ , shaded), temperature anomaly ( $^{\circ}C$ , contour, subtracting the zonal mean between  $80^{\circ}W$  and  $0^{\circ}$ ) and relative vorticity ( $s^{-1}$ , contour) at  $13^{\circ}N$  for 00z 14 Sep, across the center of the newly named Tropical Storm Helene.

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## Auxiliary Material

Two additional figures are provided to allow a comparison with Fig. 1 and Fig. 5 in the paper.

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## Caption:

Hovmoller of 850 hPa relative vorticity (left panel) and total precipitable water (right panel) from the NCEP operational analyses, latitudinally averaged (10-18N), covering the period from 1 to 15 September.

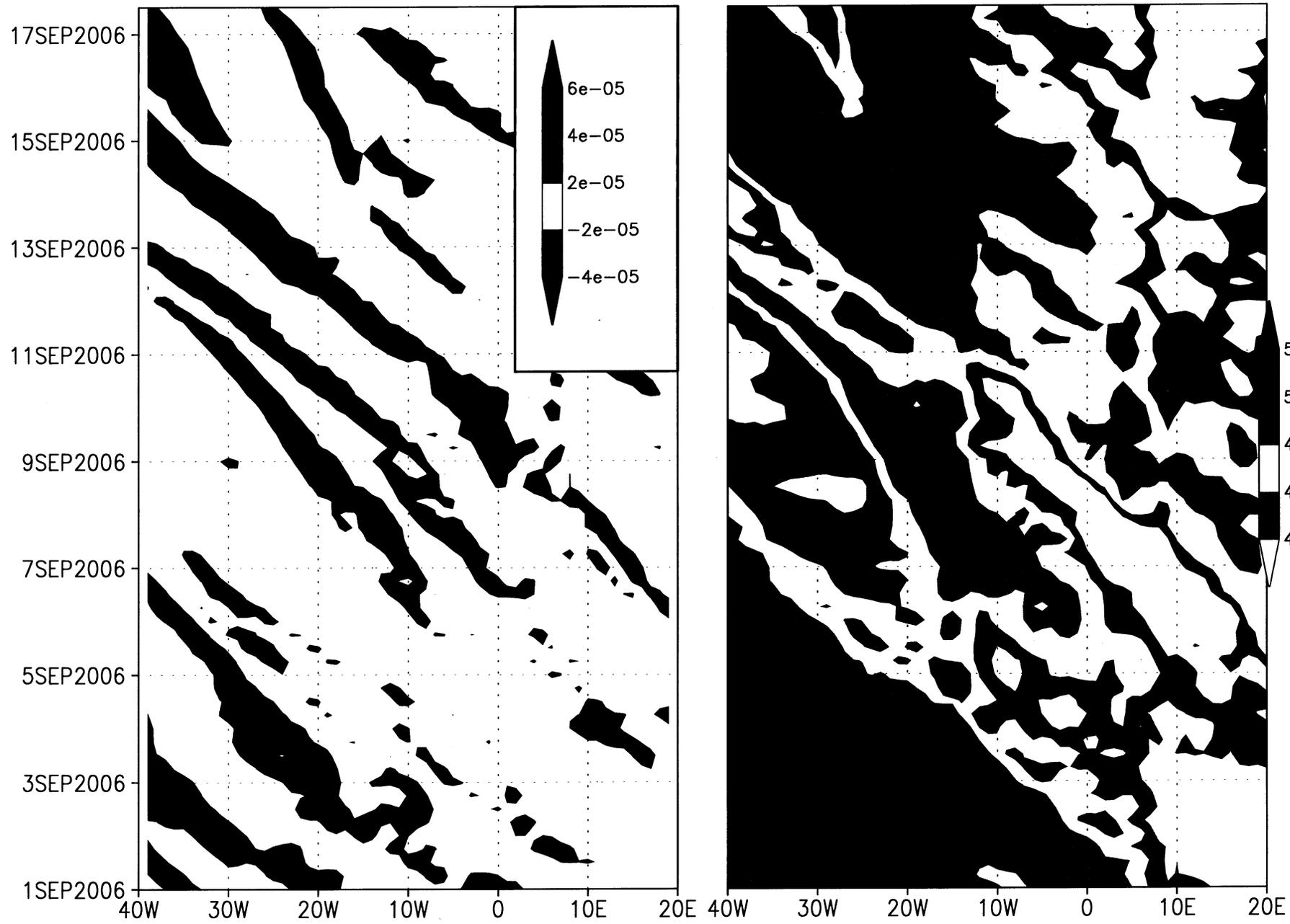
Data on pressure levels interpolated on a 1 degree resolution grid.

No significant difference from the original data in sigma levels (not shown).

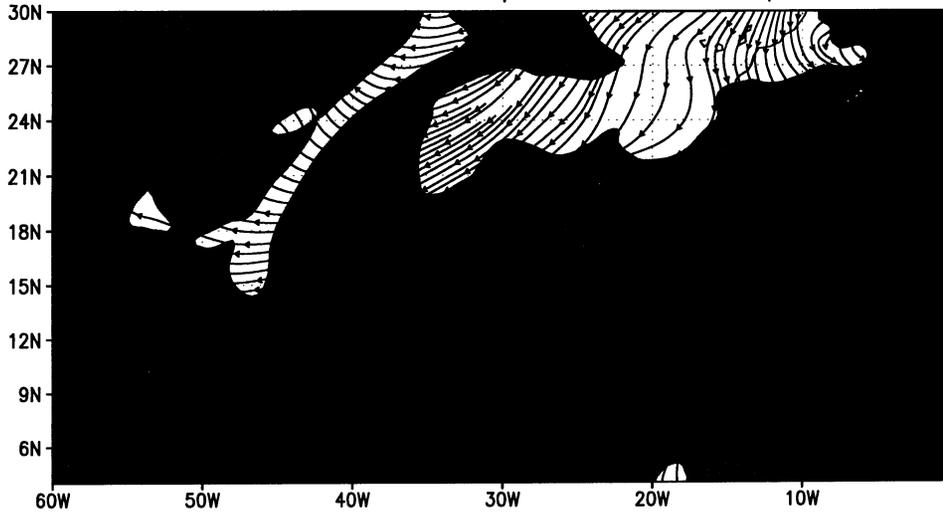
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## Caption:

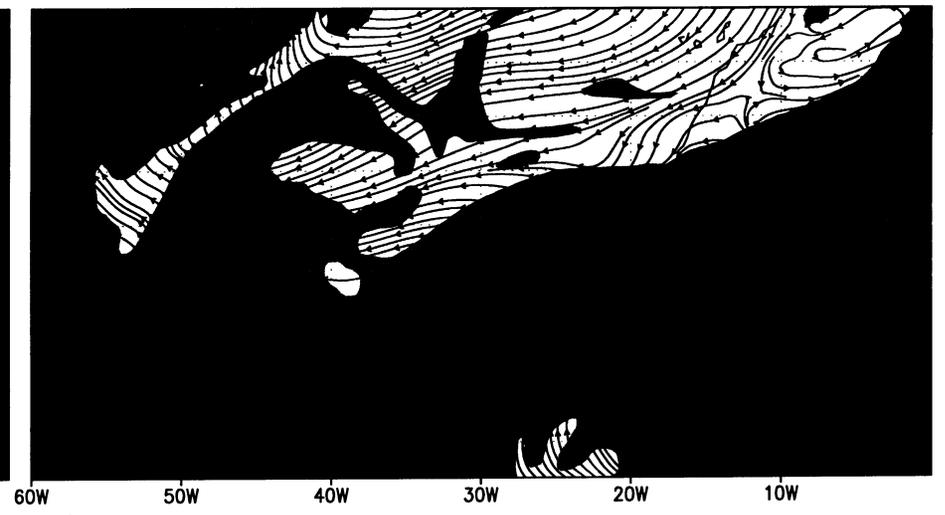
GEOS-5: 700 hPa specific humidity (Kg/Kg) and 850 hPa wind (streamlines) in the NCEP-derived initial conditions (upper left) for 00z 13 September, and relative to the 24, 48 and 72 hour forecasts for 00z 14, 15 and 16 September.



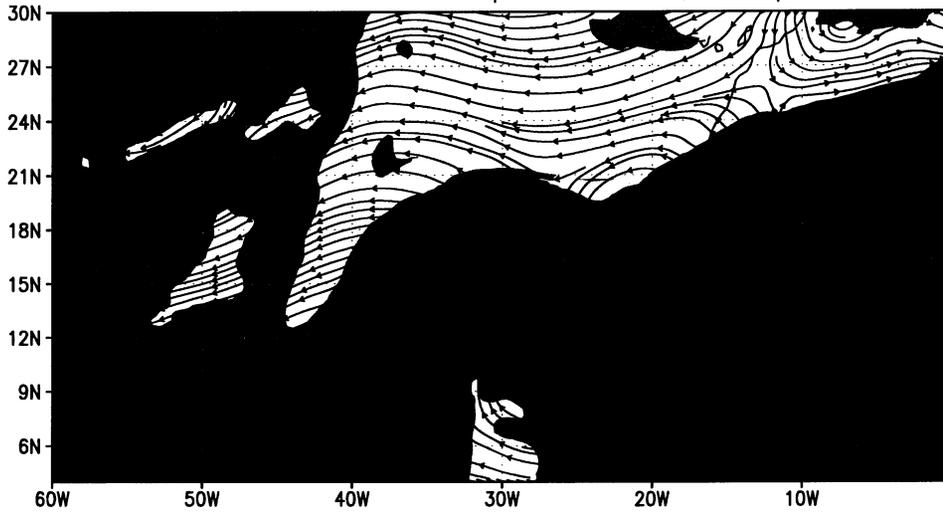
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GEOS5 IN00z13Sep06 72hF VER16Sep

