The properties of large amplitude whistler mode waves in the magnetosphere: Propagation and relationship with geomagnetic activity

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[1] We present results of a study of the characteristics of very large amplitude whistler mode waves inside the terrestrial magnetosphere at radial distances of less than 15 Re using waveform capture data from the Wind spacecraft. We observed 247 whistler mode waves with at least one electric field component (105/247 had ≥ 80 mV/m peak-to-peak amplitudes) and 66 whistler mode waves with at least one search coil magnetic field component (38/66 had ≥ 0.8 nT peak-to-peak amplitudes). Wave vectors determined from events with three magnetic field components indicate that 30/46 propagate within 20° of the ambient magnetic field, though some are more oblique (up to ∼ 50°). No relationship was observed between wave normal angle and GSM latitude. 162/247 of the large amplitude whistler mode waves were observed during magnetically active periods (AE > 200 nT). 217 out of 247 total whistler mode waves examined were observed inside the radiation belts. We present a waveform capture with the largest whistler wave magnetic field amplitude (≥ 8 nT peak-to-peak) ever reported in the radiation belts. The estimated Poynting flux magnitude associated with this wave is ≥ 300 μW/m², roughly four orders of magnitude above estimates from previous satellite measurements. Such large Poynting flux values are consistent with rapid energization of electrons. Citation: Wilson, L. B., III, C. A. Cattell, P. J. Kellogg, J. R. Wygant, K. Goetz, A. Breneman, and K. Kersten (2011), The properties of large amplitude whistler mode waves in the magnetosphere: Propagation and relationship with geomagnetic activity, Geophys. Res. Lett., 38, L17107, doi:10.1029/2011GL048671.

1. Introduction

[2] Whistler mode waves are one of the most ubiquitous wave modes observed in space plasmas. They have been observed in the magnetosphere [Burris and Helliwell, 1969], in the solar wind [Neubauer et al., 1977], upstream of interplanetary shocks [Wilson et al., 2009], upstream of planetary bow shocks [Hoppe et al., 1981], and in cometary foreshocks [Tsurutani et al., 1987]. Whistler mode waves are a right-hand polarized electromagnetic mode that can propagate along the magnetic field, obliquely to the magnetic field, or nearly perpendicular to the magnetic field as a quasi-electrostatic mode near the resonance cone [Bell and Ngo, 1990]. For simplicity, we will not distinguish between cho-

rus, plasmaspheric hiss, or oblique whistler-modes in our use of the term whistler mode wave. In the magnetosphere, whistler mode waves are thought to be driven unstable by electron temperature anisotropies [Kennel and Petschek, 1966]. Because whistler mode waves interact strongly with energetic particles [Kennel and Petschek, 1966; Lyons et al., 1972], it has been well accepted that they play an important role in global radiation belt dynamics. Thus, whistler mode waves have been a topic of extreme interest for over 40 years in magnetospheric physics.

[3] For the past 30 years plasma wave measurements have primarily consisted of time-averaged spectral intensity data. Typical time-averaged spectral zero-to-peak amplitudes for whistler mode waves are ∼0.5 mV/m for the electric field [Meredith et al., 2001] and ∼0.01–0.1 nT for the magnetic field [Horne et al., 2003, 2005]. Tsurutani et al. [2009] showed that time-averaged spectral amplitudes can severely underestimated the instantaneous wave amplitudes. This discrepancy was exemplified with the discovery of very large amplitude (>200 mV/m) whistler mode waves in the radiation belts using the STEREO spacecraft [Cattell et al., 2008], and in later observations by the Wind [Kellogg et al., 2011] and THEMIS [Cully et al., 2008] spacecraft. Cattell et al. [2008] also showed that these waves were capable of energizing electrons by more than an MeV in less than a second, which is consistent with test particle simulations [Omura et al., 2007; Bortnik et al., 2008]. Furthermore, recent observations have shown evidence that these large amplitude whistler mode waves are capable of trapping [Kellogg et al., 2010] and prompt scattering into the loss-cone of radiation belt particles [Kersten et al., 2011]. These observations have raised new questions regarding the energization and lifetime of radiation belt particles.

[4] During an eight year period, the Wind spacecraft went through a number of petal orbits within the terrestrial magnetosphere. We report on whistler mode wave observations for 13 of those petal orbits, building on the study by Kellogg et al. [2011], who used an automated search algorithm that only examined a subset of the available waveform captures to find whistler mode waves. We examined all of the waveform captures and obtained significantly more whistler mode waves. The paper is organized as follows: Section 2 introduces and outlines the data sets and analysis techniques, Section 3 describes the observations, and Section 4 discusses the conclusions of our study.

2. Data Sets and Analysis

[5] Waveform captures were obtained from the Wind/ WAVES instrument [Bougeret et al., 1995], using the time
domain sampler (TDS) receiver, which provides a waveform capture (herein called TDS sample) of 2048 points with timespans ranging from \( \sim 17 \) ms to \( \sim 1000 \) ms, depending on sample rate. The TDS receiver has both a fast (TDSF) and slow (TDSS) receiver. The TDSF receiver returns two electric field components in the spacecraft spin plane (roughly the XY/uni plane), while the TDSS receiver returns four vector components, either three electric and one magnetic or vice versa. A more detailed description of the TDS onboard memory buffer and evaluation criteria is given by Kellogg et al. [2011]. We also add that the TDS instrument continuously returns both the most recent and the best quality (typically defined by amplitude) waveform in its telemetry stream. The best quality waveforms receive more telemetry and thus bias the transmitted waveform captures to larger amplitude waves. All the amplitudes reported herein will be peak/uni to peak.

[6] We use the Wind/3DP [Lin et al., 1995] solid state telescopes (SSTs) to identify the periods when Wind entered (exited) the radiation belts, defined by the sharp increase (decrease) in omni-directional flux of the high energy electrons (SST Foil) and protons (SST Open), respectively, from the Wind/3DP instrument. At the bottom of the plot are tick mark labels of MLT, L-Shell, and radial distance along with the local UT. The vertical black lines correspond to TDS samples identified as whistler mode waves. The radiation belts are shown by the shaded region and the magnetopause crossing by the vertical blue line near 22:00 UT. The five lowest energy bins for the SST Foil and three lowest for the SST Open plots show a decrease in flux as Wind reaches perigee. This is not physical and is identified as saturation of detector at these energies [McFadden et al., 2007].

[7] The wave vector, \( \mathbf{k} \), and propagation angle with respect to the magnetic field, \( \theta_{\mathbf{k}B} \), were determined using Minimum Variance Analysis (MVA) [Khrabrov and Sonnerup, 1998] on bandpass filtered TDSS samples with three magnetic field components. The frequency ranges for each bandpass filter, determined from spectral analysis, were chosen independently for each TDS sample. We required intermediate to minimum eigenvalues of the spectral matrix, \( \lambda_{\text{mid}}/\lambda_{\text{min}} \), to satisfy the condition \( \geq 5.0 \) if less than 50 field vectors were used in the analysis [Khrabrov and Sonnerup, 1998]. All reported wave normal angles were calculated using MVA on three magnetic field component TDSS samples.

[8] We performed a search of 13 Wind petal orbits between 1998-11-13 and 2002-10-10, examining every TDS sample observed inside the magnetopause and within an L-shell of 15. We limited the results to \( L \leq 15 \) because of the small number of TDS events at higher L-shells and difficulties in the identification of a clear magnetopause crossing. Though we observed whistler mode waves inside the magnetosphere on all 13 petal orbits, only 12 orbits had whistler mode waves inside an L-shell of 15. We define a TDS sample as a whistler mode wave when it is elliptically or circularly right-hand polarized with respect to the magnetic field and has \( f_{lh} \leq f \leq f_{ce} \), where \( f_{lh} (= \sqrt{f_{ci}f_{cs}}} \) is the lower hybrid frequency and \( f_{ce} \) is the cyclotron frequency of species \( s \). We observed 247 waveform captures identified as whistler mode waves, 191 using electric fields and 56 using magnetic fields.

3. Observations

[9] Figure 1 presents an example of a Wind perigee pass from 1998-11-13/13:00:00 UT to 1998-11-14/02:00:00 UT. The vertical black lines indicate whistler mode wave
Fourier bandpass filter (frequency range shown in Figure 2, black) to remove superposed low and high frequency signals. The first panels show the $$B_y$$ (blue, instrument coordinates), second panels show $$E_x$$ (red), third panels show $$E_{x,1}$$ (green), and the fourth panels show $$E_{x,2}$$ (magenta). The fifth panels show wavelet transforms of the unfiltered $$E_{x,2}$$-component plotted on a linear frequency scale from ~700–4000 Hz, with corresponding power spectrum scale shown below [Torrence and Compo, 1998]. Note that the equatorial electron cyclotron frequencies for these two waves are ~12.94 kHz for the 1998-11-13 whistler mode wave and ~7.16 kHz for the 2000-04-10 event.

[10] Figure 2 plots two whistler mode waves observed near the equatorial plane, post midnight near $$L$$ ~4. These whistler mode waves were obtained from the TDSS instrument (with three electric and one magnetic field components sampled at 7.5 kHz). We have used the ambient magnetic field, $$B_m$$, and the sun direction, $$X_{GSE}$$, to create an orthogonal field-aligned coordinate (FAC) system. The first axis is parallel to $$B_m$$ (red, $$E_x$$ in Figure 2) and another axis is parallel to $$B_y$$ ($$|B_y| \times X_{GSE}$$ (magenta, $$E_{x,2}$$). The final axis (green, $$E_{x,1}$$) completes the right-handed basis. The waveforms were rotated into FACs and then we applied a standard Fourier bandpass filter (frequency range shown in Figure 2, top) to remove superposed low and high frequency signals. The L-shell, MLT, GSM latitude, $$f_{th}$$, and the local $$f_e$$ for each corresponding whistler mode wave are labeled in Figure 2.

[11] The example in Figure 2a, at 8 nT peak-to-peak, contains the largest whistler mode wave magnetic field ever reported in the magnetosphere. This amplitude results from only one component of the magnetic field, therefore the actual magnetic amplitude is likely to be larger. The electric field for this waveform saturated the instrument so the electric field peak-to-peak amplitude is also larger than the ~300 mV/m seen in the $$E_{x,2}$$-component. As one can see in the figure, the polarizations of the two perpendicular components are highly elliptical and significant fractions (~33%) of their electric field components are parallel to the ambient magnetic field, which implies that they propagate at oblique angles to the magnetic field. Such anomalously high field-aligned electric fields were recently reported by Breneman et al. [2011] in the inner radiation belt.

[12] Recent observations [Kellogg et al., 2010; Kersten et al., 2011] have shown that these large amplitude whistler mode waves can strongly interact with high energy electrons. Therefore, we use the relationship for the maximum change in kinetic energy, $$(\Delta KE)_{max}$$ of an electron interacting with a parallel propagating whistler derived by Omura et al. [2007, equation 44] to estimate the maximum kinetic energy these waves can impart to electrons. We find that the wave in Figure 2a could produce $$(\Delta KE)_{max} \sim 61$$ MeV electrons, assuming propagation parallel to the magnetic field. The impact of oblique propagation on the above calculation is beyond the scope of this paper but has been addressed using the test particle simulations of Roth et al. [1999] in the context of the radiation belts [Cattell et al., 2008; Kersten et al., 2011]. Although we cannot calculate the full wave Poynting flux since we only have one electric field component for the examples in Figure 2, we can estimate a lower bound on the wave Poynting flux. We found the wave in Figure 2a had as a lower bound Poynting flux magnitude of $\gtrsim 300 \mu W/m^2$, while the wave in Figure 2b had $\gtrsim 30 \mu W/m^2$. Even though these estimates are lower bounds on the total Poynting flux magnitude, they are still 3–4 orders of magnitude larger than the estimates found by Santolik et al. [2010] ($\sim 0.05 \mu W/m^2$), and are thus consistent with much more rapid energization of relativistic electrons. The examples show that: (1) large amplitude whistler mode waves in the radiation belts are bursty; (2) electric fields are in excess of two orders of magnitude above previous time-averaged spectral intensity observations consistent with those reported by Cattell et al. [2008] and Cully et al. [2008]; and (3) the search coil
magnetic field amplitudes exceed those of previous time-
averaged spectral intensity observations by two orders of
magnitude.

[15] We observed 247 whistler mode waves observed with
at least one electric field component and 66 with at least one
magnetic field component in our study. Of the 247 total
TDS samples, 217 whistler mode waves occurred within
the radiation belts. Roughly 104/191 of the whistler mode
waves with two (TDSF) or three (TDSS) electric field com-
ponents have peak-to-peak amplitudes $|\mathbf{E}| \geq 80 \text{ mV/m}$ and
28/56 of the whistler mode waves with three magnetic field
components (TDSS) have $|\mathbf{B}| \geq 0.8 \text{ nT}$. The mean plus or
minus the standard deviation of the mean of the peak-to-
peak wave amplitudes for the 191 electric field whistler
mode waves were $138 \pm 10 \text{ mV/m}$, whereas the 56 magnetic
field whistler mode waves were $1.0 \pm 0.2 \text{ nT}$. Due to the
limitations of data rate and the selection criteria for the TDS,
it is not possible to determine what percentage of the time
these very large amplitude waves occur in the radiation
belts. This question will be resolved by the upcoming RBSP
mission.

[14] We observed the waves over a broad range of MLT
and a small range of GSM latitudes, $\lambda_{\text{GSM}}$. We observed
73/247 whistler mode waves between 21 hrs $< \text{MLT} \leq
24 \text{ hrs}$ and 81/247 are observed between 0 hrs $\leq \text{MLT} \leq
3 \text{ hrs}$, the majority of which were inside 6 $R_E$ in the radi-
ation belts. We observed 114/247 waves with $|\lambda_{\text{GSM}}| \leq
5^\circ$, 52/247 with $5^\circ < |\lambda_{\text{GSM}}| \leq 10^\circ$, 22/247 with $10^\circ <
|\lambda_{\text{GSM}}| \leq 15^\circ$, 50/247 with $15^\circ < |\lambda_{\text{GSM}}| \leq 20^\circ$, and
9/247 with $|\lambda_{\text{GSM}}| > 20^\circ$. There were 46 TDSS samples with
three magnetic field components that satisfied our criteria for
the use of MVA. We found that the waves had a broad
range of wave normal angles with respect to the magnetic
field ($0^\circ \leq \theta_{\text{AB}} < 50^\circ$), where 30/46 of the waves had $\theta_{\text{AB}} \leq 20^\circ$,
10/46 had $20^\circ < \theta_{\text{AB}} \leq 30^\circ$, and 6/46 had $\theta_{\text{AB}} > 30^\circ$. For these
46 TDSS samples, we observed no relationship between $\theta_{\text{AB}}$
and GSM latitude.

[15] We observed the majority of the waves, 162/247,
during magnetically active periods with $\text{AE} > 200 \text{ nT}$. When
we examined the dependence of the wave amplitude on
one minute AE-Index estimates (not shown), we found a
slight increase in wave amplitude with increasing $\text{AE}$,
consistent with previous observations [Li et al., 2009]. A
cursory examination of LANL geosynchronous low energy
(30–300 keV) electron injections show large injections on
the nightside just prior to the two whistler mode waves observed
in Figure 2. Both the LANL observations of particle injec-
tions and the increasing wave amplitude with increasing AE
would be consistent with increased substorm activity pro-
viding free energy for wave growth. We did observe some
evidence of substorm injections in the LANL data for every
orbit examined herein. However, as discussed above, we
observe as nearly as many waves on the dusk side of Earth
as the dawn. The observation of waves on the dusk-side is
not consistent with the standard picture of wave generation
by $\sim 10–100 \text{ keV}$ electrons injected near midnight and drift-
ing around towards dawn [Thomsen, 2004].

4. Discussion and Conclusions

[16] We present a study of the properties of large ampli-
tude whistler mode waves in the radiation belts and outer
magnetosphere to provide a more detailed characterization
of large amplitude whistler mode waves described by Cattell
et al. [2008], Cully et al. [2008], and Kellogg et al. [2011].
We have also presented the largest whistler wave magnetic
field ($\sim 8 \text{ nT}$ peak-to-peak) observations ever reported in
the radiation belts, with a corresponding Poynting flux magni-
tude $\sim 300 \mu \text{W/m}^2$, nearly four orders of magnitude above
previous time-averaged spectral intensity observations. Mini-
mum variance analysis showed that these large amplitude
whistler mode waves can propagate at very oblique angles
(up to $\sim 50^\circ$) with respect to the ambient magnetic field, but
30/46 propagate within $20^\circ$.

[17] The majority of the whistler mode waves in our study
were observed with frequencies less than or equal to half the
local (and equatorial) electron cyclotron frequency with a
mean frequency of $\sim 1.4 \pm 1.0 \text{ kHz}$. The mean value of the
wave frequency normalized by the local(equatorial) electron
cyclotron frequency is $\sim 0.25 \pm 0.17(\sim 0.27 \pm 0.26)$. We
observed 114/247 waves with $|\lambda_{\text{GSM}}| \leq 5^\circ$, 166/247
with $|\lambda_{\text{GSM}}| \leq 10^\circ$, and 188/247 with $|\lambda_{\text{GSM}}| < 15^\circ$. Of the
247 whistler mode waves, 217 were observed in the radia-
tion belts. We found 154/247 of the whistler mode waves to
be located between 21 hrs $\leq \text{MLT} \leq 3 \text{ hrs}$ (i.e., nightside)
while 39/247 were between 6 hrs $\leq \text{MLT} \leq 18 \text{ hrs}$ (i.e.,
dayside). These waves have large peak-to-peak amplitudes
(92/247 have $\geq 100 \text{ mV/m}$ and 34/66 have $\geq 1.0 \text{ nT}$). Nearly
all of the waves were observed in association with LANL low
energy (30–300 keV) electron injections and 162/247 were
observed during magnetically active periods ($\text{AE} > 200 \text{ nT}$).
Furthermore, the wave amplitudes show an increasing
trend with increasing $\text{AE}$, consistent with increased sub-
storm activity providing more free energy for whistler growth
and previous observations [e.g., Li et al., 2009]. However,
we observed roughly an equal number of waves between
21 hrs $< \text{MLT} \leq 24 \text{ hrs}$ (73/247) and 0 hrs $\leq \text{MLT} \leq 3 \text{ hrs}
(81/247), suggesting some of the waves could not be caused
by the standard model of substorm injection [Thomsen,
2004]. Note that the limited coverage of the Wind perigee
passes makes it impossible to determine the MLT distribution
of the waves in a statistically meaningful way.

[18] Only recently with the introduction of waveform cap-
ture instruments have we begun to understand the range of
amplitudes of whistler mode waves in the radiation belts.
The largest amplitude waves were observed between 0 hrs $\leq
\text{MLT} \leq 6 \text{ hrs}$ (i.e., dawn side), inside 6 $R_E$ in the radia-
tion belts, and within $\pm 10^\circ$ of the magnetic equator. We will
illustrate the relative distribution of events returned by the TDS
instrument by examining the two perigee passes that contain
the examples shown in Figure 2, where Wind observed
some of the largest whistler events ever recorded in the mag-
netsphere. Wind observed 11 TDSF and 3 TDSS samples
consistent with whistler mode waves (in roughly 6 minutes)
for the 1998–11–13 perigee pass, whereas Wind observed
27 TDSF and 7 TDSS samples (in roughly 1.75 hours) for
the 2000–04–10 perigee pass. All 14 TDS samples for the
1998–11–13 perigee pass had peak-to-peak amplitudes $|\mathbf{E}| \geq
200 \text{ mV/m}$ and all 3 TDSS samples had $|\mathbf{B}| \geq 8 \text{ nT}$. Of the
34 TDS samples observed in the 2000–04–10 perigee pass,
31 had peak-to-peak amplitudes $|\mathbf{E}| \geq 160 \text{ mV/m}$ and all
7 TDSS samples had $|\mathbf{B}| \geq 2 \text{ nT}$. These results are consistent
with STEREO [Cattell et al., 2008; Breneman et al., 2011]
and THEMIS [Cully et al., 2008] studies. Cattell et al. [2008]
observed 24 waveform captures with whistler mode waves in
$\sim 4 \text{ minutes}$ all with peak-to-peak amplitudes $\geq 200 \text{ mV/m}$. 

4 of 5


