Mean dynamic topography of the Arctic Ocean

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1. Introduction

Satellite laser and radar altimeters on ICESat and Envisat provide measurements of both sea ice elevation and sea surface height (SSH) across the Arctic Ocean, using lead and open water elevation within the sea ice pack. First, these data were used to derive two independent mean sea surface (MSS) models by stacking and averaging along-track SSH profiles gathered between 2003 and 2009. The ICESat and Envisat MSS data were combined to construct the high-resolution ICEn MSS. Second, we estimate the 5.5-year mean dynamic topography (MDT) of the Arctic Ocean by differencing the ICEn MSS with the new GOCO02S geoid model, derived from GRACE and GOCE gravity. Using these satellite-only data we map the major features of Arctic Ocean dynamical height that are consistent with in situ observations, including the topographical highs and lows of the Beaufort and Greenland Gyres, respectively. Smaller-scale MDT structures remain largely unresolved due to uncertainties in the geoid at short wavelengths. Citation: Farrell, S. L., D. C. McAdoo, S. W. Laxon, H. J. Zwally, D. Yi, A. Ridout, and K. Giles (2012), Mean dynamic topography of the Arctic Ocean, Geophys. Res. Lett., 39, L01601, doi:10.1029/2011GL050052.

2. Satellite Data

2.1. ICESat Laser Altimetry

We utilize ICESat laser altimetry data from the October-November, February-March, and May-June observation periods, gathered over 5.5 years between October 2003 and April 2009 (see Yi et al. [2011] for specific laser operation dates). We exclude from our analysis data collected during the first and last campaigns, due to uncorrected range errors in the former, and the short 12-day time-span of the latter. The data are Release 428 processed using precision orbit, tidal, saturation and inverse barometer corrections with cloud-filtering techniques used to exclude data corrupted by forward scattering [Zwally et al., 2008; Yi et al., 2011]. Instantaneous SSH was calculated at the full (40 Hz) along-track resolution following the methodology of Zwally et al. [2008] and based on the assumption that the small laser footprint can distinguish thin ice and narrow leads, where the precision of lead elevations is ~2–3 cm [Kwok and Morison, 2011]. A MSS profile is computed for each reference ground-track, by stacking and averaging 16 repeat passes gathered during the 5.5-year observation period. At least seven SSH measurements were acquired for the majority of reference footprints, mitigating the impact of inter-campaign biases on the 5.5-year average. The dense along-track sampling (~50 m every 170 m) provided by ICESat is ideal for deriving MSS profiles even across steep sea surface slopes at continental shelf-breaks and over narrow bathymetric features.
gravity field than GRACE and has already improved global gravity in the wavelength band ranging from spherical harmonic degree 100 to 250 (~160–400 km) [Pail et al., 2010]. Here we use the GOCO02S geoid, a geopotential model derived from the latest satellite gravity based on 8 months of GOCE data, 7 years of GRACE data, GPS tracking data, and CHAMP data [Goiginger et al., 2011]. In contrast to the Arctic Gravity Project (ArcGP) [Kenyon and Forsberg, 2008] and Earth Gravitational Model 2008 (EGM2008) [Pavlis et al., 2008] geoids, which assimilate satellite altimetry, GOCO02S is a purely gravitational model. At wavelengths greater than 500 km the accuracy of satellite gravimetric geoid is comparable to altimetry measurements, at ~2 cm [Pail et al., 2010].

3. The ICEn MSS

Measuring SSH in ICESat data over the open ocean can be affected by long-wavelength ocean waves. Therefore, the ICESat MSS is computed for the ice-covered Arctic Ocean only. Envisat MSS measurements are available over both the ice-covered and open ocean, thereby providing coverage south to 66 °N. To eliminate any seasonal signals not captured during ICESat’s intermittent measurement campaigns, a subset of Envisat data were selected for temporal coincidence with ICESat observations. The combined ICESat-Envisat MSS, the ICEn MSS of the Arctic Ocean, was constructed as follows: Envisat data were used exclusively south of the ice edge to 66 °N, ICESat and Envisat data were used north of the ice edge to 81.5 °N, ICESat data were used exclusively above 81.5 °N to 86 °N. Individual along-track MSS height profiles were combined on a grid with longitude spacing of 1/20° and latitude spacing of 1/50° using a weighted “nearest-neighbor” interpolation scheme to create the ICEn MSS. The ICEn MSS provides continuous mapping of basin-scale Arctic Ocean topography to 86 °N (Figure 1), and updates and extends the first mapping of the Arctic MSS [Peacock and Laxon, 2004]. The variability of the MSS reflects both sea floor topography and density variations in the oceanic crust and upper mantle [Peacock and Laxon, 2004; Laxon and McAdoo, 1994]. Artificial illumination of the surface (Figure 1) reveals that the ICEn MSS resolves the steepest sea surface slopes associated with bathymetrically prominent features such as the Mohns-Knipovich, Lomonosov and Gakkel Ridges, and the shelf-breaks off Eastern Greenland and north of the Canadian Arctic Archipelago.

4. Assessment of Satellite Altimeter MSS Models

We assess the spatial resolution of the ICESat and Envisat MSS models via comparison with the recently published DNSC08 MSS, a global field that provided the first Arctic MSS map complete to 90 °N [Andersen and Knudsen, 2009]. The DNSC08 MSS combines satellite radar (ERS-2 and Envisat) and laser (ICESat) altimetry, with the ArcGP.06 geoid beyond 86 °N [Kenyon and Forsberg, 2008].

We conduct statistical analyses of the ICESat-only, Envisat-only, and DNSC MSS models in “Region A”, an area bounded by the coordinates 116 °E–238 °E, 76.8 °N–81.4 °N (white box, Figure 1). This represents the largest area in the central Arctic that is uncontaminated by island chains.
and sampled by both ICESat and Envisat. There is excellent agreement in the mean difference (5 cm) and standard deviation (2 cm) between the ICESat and Envisat MSS models in Region A (not shown). The comparison between either the ICESat or Envisat MSS fields and the DNSC08 MSS reveals a larger mean difference, with an offset of 19–21 cm and a standard deviation of 7–8 cm. This offset may be due to the treatment of altimetry data in the construction of the DNSC08 MSS at high northern latitudes, where Andersen and Knudsen [2009] use alternative methodologies to estimate altimetric SSH and apply a number of offset corrections.

[11] Coherency between the three MSS models in Region A is above 0.5 at wavelengths greater than ~30 km (Figure 2a), indicating agreement in the expression of topographical features wider than ~30 km. At shorter wavelengths however there is higher coherency between the ICESat and Envisat MSS fields, than between the ICESat and DNSC08 fields (Figure 2a). While the ICESat MSS and Envisat MSS use full-resolution along-track data (40 Hz and 20 Hz respectively), 1 Hz (i.e., 1 s) SSH observations were used in the construction of the DNSC08 MSS Andersen and Knudsen [2009]. Short wavelength features are therefore attenuated in DNSC08 and could contribute to the lower coherency observed at short wavelengths (Figure 2a). We also calculate the power spectrum of the differences between the models in Region A, to examine noise in the MSS fields in the wavelength band 15–1000 km (Figure 2b). Superior agreement (lower noise) between ICESat and Envisat, than with the DNSC08 MSS is indicated, particularly at wavelengths >75 km. Residual orbit errors in the ICESat and Envisat data (appearing as ground-track striations in the MSS fields) may contribute to higher noise levels at shorter wavelengths.

[12] Taken together the results show greater agreement between the satellite-only Arctic Ocean MSS models derived independently from ICESat and Envisat altimetry, with no overall bias between these fields, compared to the current state-of-the-art (e.g., DNSC08). This demonstrates that by retaining the high-resolution along-track data provided by satellite altimeters, the expression of even extremely narrow topographical features can be resolved.

5. Mean Dynamic Topography

[13] Basin-scale estimates of MDT are vital for estimating ocean structure at both short and long wavelengths for ocean circulation studies and validation of geostrophic current velocity from models. Arctic Ocean MDT remains poorly constrained due to a lack of widespread and routine observational data. The MDT, \( h_{MDT} \), reported with respect to a temporal averaging period, may be determined as the difference between the altimetric MSS height (\( h_{MSS} \)) and the geoid (\( h_G \)): \( h_{MDT} = h_{MSS} - h_G \), where the effects of tides, atmospheric pressure fluctuations, and sea ice elevation have been removed from \( h_{MSS} \). Knowledge of \( h_G \) to an accuracy of a decimeter or better is required to accurately derive MDT across the Arctic Basin. We compute two independent estimates of the long-term, 5.5-year, Arctic MDT spanning 2003 to 2009. The estimates represent the longest-term measurement of Arctic MDT currently feasible using ICESat data and temporally coincident Envisat data. Our approach reduces the effects of seasonal signals and inter-campaign bias, or other altimetric range errors, and thus represents an improvement over previous MDT mappings using subsets of the ICESat data record [e.g., Kwok and Morison, 2011; Forsberg et al., 2007]. First we difference the high-resolution ICEn MSS with the EGM2008 geoid (Figure 3a). We compare this preliminary estimate of Arctic MDT to a second, satellite-only estimate computed relative to the GOCO02S geoid (Figure 3b). Attenuation of short-wavelength gravity signals is inversely proportional to the altitude of GOCE (~260 km) and necessitates some filtering of the MDT field. Filtering also reduces residual noise in the MSS field due to unmodeled tidal or atmospheric effects, ground-track striation, etc. Here we filter both MDT fields using an isotropic Gaussian filter with a 250 km half-wavelength. A zonal MDT profile across the Arctic Basin at 83 °N, between 135 °E and 250 °E (Figure 3c) illustrates the value of filtering in reducing noise. There are notable differences
between the two MDT estimates including an overall offset of 9 cm across the Arctic Basin.

Our MDT maps (Figures 3a and 3b) reveal the primary large- and meso-scale steady-state circulation features of the Arctic Ocean, and are consistent with the mean dynamical height climatology, derived from in situ observations [Arctic Climatology Project, 1997], including the Beaufort Gyre, a topographical high of ~0.35 m centered at ~215 °E, 74 °N, and a ~0.85 m gradient in the dynamic topography across the Arctic Basin to the Greenland Gyre, a low of ~0.5 m centered at 0 °E, 75 °N in the northern Greenland Sea. Additional circulation features including the East Greenland Current and Transpolar Drift are resolved, particularly in the satellite-only estimate (Figure 3b).

Figure 3a reveals anomalies in the Canada Basin (at 240 °E, 79 °N) and Lincoln Sea (83 °N, 300 °E) as well as meandering MDT contours in the eastern Arctic (82–85 °N, 60–175 °E). Remaining uncertainties in the EGM2008 geoid are the likely contributor to these anomalies. Figure 3c reveals noise in the MDT field derived using EGM2008, particularly at 150 °E, 165 °E, 181 °E, and 230 °E, that is not present in the MDT derived using GOCO2S. EGM2008 is derived from a combination of airborne, surface and submarine gravimetry with GRACE data at long wavelengths and satellite altimetry over the oceans. While EGM2008 yields a highly accurate geoid (to ~0.06 m) over the major ocean basins [Pavlis et al., 2008], geoid errors are substantially larger in the Arctic, particularly short-wavelength errors of several decimeters to 1 m (D. C. McAdoo et al., Gravity of the Arctic Ocean from satellite data with validations using airborne gravimetry, manuscript in preparation, 2011), where ArcGP is used as a primary input. Furthermore, since EGM2008 assimilates altimeter data in the Arctic Ocean it is hence contaminated by residual ocean signals. GOCO02S, derived entirely from satellite gravity data, is free from sea surface topography signals and thus provides an improved approach for accurate derivation of Arctic MDT on basin scales [Tapley et al., 2003].

6. Summary

We have constructed a new MSS model of the Arctic Ocean, the ICEn MSS, derived from a combination of 5.5 years of ICESat-laser and Envisat-radar altimetry. The ICEn MSS mapping has high spatial resolution and contains details of the steepest sloping sea surface topography, representing a significant improvement over comparative MSS models such as the DNSC08 MSS. Indeed due to the remaining short-wavelength errors in sectors of the best available Arctic geoids, high-resolution altimetric MSS models, such as the ICEn MSS, provide a useful reference surface for more accurate derivation of sea ice freeboard from CryoSat-2, IceBridge and ICESat-2 altimetry. Additionally, along-track sea surface slopes will be used to improve models of the Arctic marine gravity field.

Differencing the high-resolution ICEn altimetric MSS and the recently released GOCO02S geoid, provides for the first time a new, satellite-only estimate of Arctic Ocean MDT filtered to 250 km. Filtering of the MDT remains necessary to remove residual noise in both the MSS and geoid fields, and thereby restricts the wavelength at which features may be resolved. We compared the satellite-only MDT to an independent estimate derived using the EGM2008 geoid. The GOCO02S geoid, constructed entirely from satellite gravity and free from sea surface topography signals, provides a superior approach for accurate derivation of Arctic Ocean MDT on basin scales, in contrast to EGM2008, which assimilates altimeter data. Our satellite-only MDT estimate describes the major features of Arctic Ocean dynamical
height including the Beaufort and Greenland Gyres, the Transpolar Drift and East Greenland Current, and reveals a steep MSS gradient from the Pacific to Atlantic sectors of the Arctic. Resolving the more detailed, small-scale structure of Arctic Ocean circulation such as stationary eddies remains elusive; the paucity of short-wavelength detail in the geoid has now become the limiting factor.

[18] The use of satellite-only data for deriving MDT demonstrates that long-term observation of Arctic Ocean circulation is now possible. Satellite radar altimeters, such as CryoSat-2, continually profile the Arctic Ocean throughout the year, such that analysis of the full, long-term radar altimeter dataset should provide details of seasonal SSH variability. The ICEn MSS offers complete coverage from 66 °N to 86 °N; new observations from CryoSat-2 should help to fill the “polar hole” above 86 °N to 88 °N. Observational continuity into the latter half of this decade is expected with the launch of ICESat-2 planned for 2016. Together with future advances in Arctic geoid modeling these satellite altimetry data will improve our understanding of Arctic Ocean circulation and its inter-annual variability.

[19] Acknowledgments. We acknowledge NASA’s ICESat Science Project and the NSIDC for distribution of the ICESat data (see http:// nsidc.org/data/icesat/), and ESA for the distribution of Envisat data. We thank three anonymous reviewers for their comments and suggestions, which helped to improve this note. Support for SLF has been provided by the NASA Cryosphere Program under NASA grant NNX10AG17G. Support for AR, KAG, and SWL is provided by the United Kingdom Natural Environment Research Council. The views, opinions, and findings contained in this report are those of the authors and should not be construed as an official National Oceanic and Atmospheric Administration or U.S. Government position, policy, or decision.

[20] The Editor thanks two anonymous reviewers for their assistance in evaluating this paper.

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