KETENE FORMATION IN INTERSTELLAR ICES: A LABORATORY STUDY

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ABSTRACT

The formation of ketene (H₂CCO, ethenone) in polar and apolar ices was studied with in situ 0.8 MeV proton irradiation, far-UV photolysis, and infrared spectroscopic analyses at 10–20 K. Using isotopically enriched reagents, unequivocal evidence was obtained for ketene synthesis in H₂O-rich and CO₂-rich ices, and several reaction products were identified. Results from scavenging experiments suggested that ketene was formed by free-radical pathways, as opposed to acid–base processes or redox reactions. Finally, we use our results to draw conclusions about the formation and stability of ketene in the interstellar medium.

Key words: astrobiology – astrochemistry – ISM: molecules – molecular data – molecular processes

1. INTRODUCTION

Ketene (H₂CCO), also called ethenone, is a 16-valence electron molecule with a photo- and thermal chemistry that has been extensively studied by laboratory scientists. Ketene readily converts into other compounds, for example combining with H₂O to make acetic acid, with NH₃ to make acetamide, and with CH₃OH to make methyl acetate. See Figure 1. Reactions of ketene and similar molecules include dimerization, acylation, cycloadditions, and polymerization. Ketene dissociation gives methylene (CH₂), which undergoes a wide variety of organic reactions, some of which give ketene an astrobiological importance. For example, CH₂ in its 3B₁ electronic ground state can insert itself into C–H chemical bonds and extend the length of a carbon chain. In its 1A₁ singlet excited state, CH₂ can add across double bonds to form cyclic compounds, and even insert into rings to enlarge them. An introduction to the extensive chemical literature on ketene is provided by reviews of Tidwell (1990, 2005) and Allen & Tidwell (2012), with older publications still being valuable (e.g., Wilmore & Stewart 1907; Staudeinger 1912; Hagemeyer 1949; Quadbeck 1956). Recent papers on synthetic applications of ketene-type molecules include Orr & Calter (2003) and Paull et al. (2009).

For astronomers, ketene has long been recognized as an interstellar molecule. It first was detected in the gas phase through the microwave measurements of Turner (1977), which were confirmed and extended by other investigators (Johansson et al. 1984; Irvine et al. 1989; Turner et al. 1999). Continued interest in this molecule is shown by recent gas-phase detections of ketene in a prestellar core (Bacmann et al. 2012, abundance ~10¹³ cm⁻²) and in an extragalactic source (Muller et al. 2011, abundance ~10⁻¹⁰ relative to H₂). The formation reactions for interstellar ketene are at present unknown and have even been said to be “subject to controversy” (Muller et al. 2011). Gas-phase production models (Herbst & Leung 1986) have been invoked, as have models in which ketene is formed on grains (Guennoun et al. 2005), and diazoacetalddehyde (Krantz 1973) all undergo solid-phase, low-temperature decomposition to give ketene, but none of these precursor molecules are expected to have a high abundance in astronomical ices.

Beyond the astrochemical literature, numerous investigations have resulted in the identification of ketene as a solid-phase reaction product, but to our knowledge all such experiments have involved chemical precursors, icer matrices, or ice components that are unlikely to occur or are of low abundance in an astronomical setting. As examples, acetone (Zhang et al. 1997), acetic acid (Macoas et al. 2004), thioacetic acid (Romano et al. 2002), acetyl chloride (Rowland & Hess 1997), acetyl cyanide (Guennoun et al. 2005), and diazoacetalddehyde (Krantz 1973) all undergo solid-phase, low-temperature decomposition to give ketene. However, none of these precursor molecules are expected to have a high abundance in astronomical ices.

Because of this paucity of laboratory data on reaction paths leading to ketene in astronomical ices, here we bring together some earlier experimental observations related to this molecule and place them, and some of our new experimental results, into a context that emphasizes the ease of ketene formation in interstellar ices. We examine two types of ices, one made of apolar components and the other dominated by H₂O, a strongly polar molecule, with an emphasis on radiation-chemical methods and support from photochemistry experiments. The present work differs from previous studies in that for ketene precursors we only use molecules that have been detected in the interstellar...
medium or that are suspected to interstellar. In this way, we con-
tinue our laboratory investigations of low-temperature reactions
that lead to astronomical molecules either seen or suspected in
interstellar and planetary environments.

2. ICES AND RADIATION CHEMISTRY

In our experiments, ices were exposed to a beam of 0.8 MeV
protons ($p^+$) to initiate radiation-chemical changes at 10–20 K.
At this energy, a single $p^+$ passing through an ice produces a
trail of thousands of ionizations and excitations as its energy
is degraded. Many such ionizations will produce secondary
electrons which, in turn, create separate tracks of yet more
ionizations and excitations leading to further chemical changes.
In this way, which we emphasize, the direct chemical action of
the initial MeV radiation is completely overshadowed by that
from secondary electrons of energies less than $\sim 20$ eV (Pimblott
& LaVerne 2007). Therefore, to a strong first approximation
the products of various ionizing radiations (e.g., $e^-$, $p^+$, He$^+$,$^\gamma$-rays) are identical, and strongly resemble those of
UV photochemistry (Spinks & Woods 1964; Hudson et al. 2001;
Hudson & Moore 2000). In the present work, this similarity in
chemical products meant that we were able to use photochemical
experiments, which were relatively fast and inexpensive, as a
way to anticipate the results of ion irradiations, which were
more costly to obtain.

A final point concerns the GeV particle energies available in
the interstellar medium in comparison to our 0.8 MeV protons.
Although GeV ions penetrate more deeply into ices than MeV
ions, the energy delivered by the former per unit distance
(stopping power) is several orders of magnitude smaller than
for MeV ions of the same type. Combined with the fact that
GeV cosmic rays are of much lower abundance than those of
smaller energy, the conclusion is that the radiation chemistry
caused by MeV $p^+$ dominates that of GeV $p^+$ in interstellar ices.

3. EXPERIMENTAL

Most of the details of our experimental procedures are already
in print, and so we present only a summary here. See Loeffler &
Hudson (2012), Ferrante et al. (2008), Hudson & Moore (2003),
and Gerakines et al. (2000) for additional information.

All chemical reagents used for experiments were commercial-
ly available and purchased from Sigma–Aldrich, Matheson,
and Cambridge Isotopes. These reagents included O$_2$, CO$_2$, CO,
CH$_4$, C$_2$H$_2$, and C$_2$H$_4$, and selected $^{13}$C and $^{18}$O isotopologues.
Slush baths were employed to remove acetone from commer-
cial acetylene. The water used to make H$_2$O-rich ices was triply
distilled and had a resistivity greater than $10^7 \Omega$ cm. Ice samples
were prepared by vapor-phase condensation for about 1 hr onto
pre-cooled (10–20 K) metal substrates (aluminum and gold) in
a vacuum chamber ($\sim 1 \times 10^{-7}$ Torr). The resulting ice thick-
nesses were 2.0–3.0 $\mu$m, and these thicknesses were measured
with a laser interference system.

Table 1 gives the ice mixtures studied in this work. These are
not strictly analogous to the compositions of some interstellar
ices, which are known to be dominated by H$_2$O or CO$_2$,
but rather were selected to probe reactions between specific
molecules, and to test for the possibility of solid-phase ketene
formation. To a first approximation, the yields we observed, and
which we report relative to the hydrocarbon component present,
will be proportional to the initial amount of hydrocarbon present
in each ice. The precise scaling law we leave for future studies.

Ices were irradiated with 0.8 MeV protons from a Van de
Graaff accelerator to doses of 1–25 eV per 16 amu molecule. For
the use of proton radiolysis to mimic cosmic-ray bombardment
see Moore et al. (2001). Since sample thicknesses were on the
order of a few micrometers, the incident protons, having a range
of 22 $\mu$m (Ziegler et al. 1985), penetrated the ices and came to
rest in the underlying metal substrate, not in the ices themselves.

Far-UV photolyses were performed with a microwave-discharge

![Figure 1](https://example.com/figure1.png)

Figure 1. Four laboratory addition reactions of ketene. See Staudinger (1912), Wang & Schueler (1949), Quadbeck (1956), and references therein for details.
Table 1

Two-Component Ice Mixtures Studied

<table>
<thead>
<tr>
<th>Mixture</th>
<th>Approximation</th>
<th>Composition</th>
</tr>
</thead>
<tbody>
<tr>
<td>CO2 + $^{13}$C2H2</td>
<td></td>
<td>1:1</td>
</tr>
<tr>
<td>O2 + C2H2</td>
<td></td>
<td>1:1</td>
</tr>
<tr>
<td>O2 + $^{13}$C2H2</td>
<td></td>
<td>1:1</td>
</tr>
<tr>
<td>$^{18}$O2 + C2H2</td>
<td></td>
<td>1:1</td>
</tr>
<tr>
<td>H2O + $^{13}$C2H2</td>
<td></td>
<td>5:1</td>
</tr>
<tr>
<td>O2 + C2H4</td>
<td></td>
<td>1:10</td>
</tr>
<tr>
<td>$^{18}$O2 + C2H4</td>
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<td>1:10</td>
</tr>
<tr>
<td>C$^{18}$O2 + C2H4</td>
<td></td>
<td>1:1</td>
</tr>
<tr>
<td>CO + CO2</td>
<td></td>
<td>1:1</td>
</tr>
<tr>
<td>$^{13}$CO2 + $^{13}$CH4</td>
<td></td>
<td>1:1, 10:1</td>
</tr>
<tr>
<td>CO + CH4</td>
<td></td>
<td>1:1</td>
</tr>
<tr>
<td>CO + $^{13}$CH4</td>
<td></td>
<td>1:1</td>
</tr>
</tbody>
</table>

hydrogen-flow lamp producing primarily Lyman photons (10.2 eV, 121.6 nm, $\sim 3 \times 10^{14}$ photons cm$^{-2}$ s$^{-1}$). See Gerakines et al. (2000) for more on the UV technique. The lamp employed was the same as the type used by Ehrenfreund et al. (2001), but with $\sim 30\%$ of their photon flux at the sample’s surface.

Infrared (IR) spectra of H$_2$O-rich ices were recorded as 60 scan accumulations from 4000 to 400 cm$^{-1}$ before and after exposure to the proton beam or UV lamp, typically at a resolution of 4 cm$^{-1}$. The inherent widths of IR absorptions in H$_2$O-dominated ices usually made it unnecessary to employ higher resolutions, but such was not the case with apolar ices. Spectra of ices made only of non-polar components (e.g., O$_2$ + C$_2$H$_2$, CO$_2$ + C$_2$H$_2$) had sharper lines and were studied at 2 cm$^{-1}$ resolution with 100 scan accumulations or, in a few cases, at 1 cm$^{-1}$ resolution. In all cases, the spectrum of an ice was divided by that of the mirror substrate, recorded before the sample’s deposition, and the resulting ratio then was converted into absorbance in the usual way ($A = -\log (I/I_0)$). Absorbance bands were integrated after subtraction of baselines that best matched the continuum on each side of the IR feature of interest.

We estimated ketene’s abundance in our irradiated ices by comparing the IR band areas of $^{13}$C-labeled ketene (2071 cm$^{-1}$) and $^{18}$O$_2$ (2281 cm$^{-1}$) in our CO$_2$ + $^{13}$C$_2$H$_2$ mixtures. For this estimate, we assumed (1) that the ratio of absolute intensities of CO$_2$ and ketene, as measured in IR transmission experiments, was the same as in our reflectance experiments, (2) that the $^{13}$CO$_2$-to-$^{18}$O$_2$ abundance ratio ($^{13}$C/$^{12}$C) was 0.011 in accord with each isotope’s natural abundance, and (3) that the band strengths for $^{13}$C and $^{12}$C isotopologues were essentially the same. The first assumption is likely the main source of error as it has long been known that IR reflectance and transmittance spectra of thin samples ($\sim 1 \mu$m) can differ and that Beer’s Law can fail (e.g., Maeda & Schatz 1961; Pacansky & England 1986). To estimate the error involved in assumption (1) we grew a CO + CO$_2$ (1:1) ice and measured the $^{13}$CO-to-$^{13}$CO$_2$ ratio, finding it to differ by as much as a factor of three from expectations (Yamada & Person 1964; Jiang et al. 1975). This led us to estimate our ketene yield (vide infra) as accurate to a factor of three, an uncertainty we hope to reduce considerably in the future. Finally, we note that for our other ices that are similar in thickness to our CO$_2$-rich samples, we also can estimate ketene’s abundance since the absolute intensities of ketene in the two cases will be similar.

To aid in assignments of spectral features, the Spartan software package of Wavefunction, Inc. was used for density-functional theory (DFT) calculations at the B3LYP/6-31+G* level to determine isotopic shifts of the most-intense vibrational bands of selected ketene and CO isotopologues. In many cases, relevant spectral data are in the literature, but our DFT calculations gave a single consistent approach to estimating shifts due to $^{13}$C and $^{18}$O labels.

4. RESULTS

Our initial interest in ketene, aside from astronomical considerations, was motivated by an experiment showing that proton-irradiated H$_2$O + C$_2$H$_2$ ices produced CO (Moore & Hudson 1998). That a set of reactions would occur in an ice to give H–C≡C–H → C≡O, converting a carbon–carbon triple bond into a carbon–oxygen triple bond, was contrary to both experience and expectations, as were similar observations in a later ice-photocatalysis experiment of Wu et al. (2002). However, the C$_2$H$_2$ to CO conversion could be understood if ketene formed in an intermediate step and then dissociated to give CO. The correct reaction sequence thus was assumed to be H–C≡C–H → H$_2$C=C=O → C≡O. Unfortunately, ketene’s strongest IR feature, a C≡C=O asymmetric stretch, is almost coincident with the intense fundamental vibration of solid CO near 2136 cm$^{-1}$, making it difficult to observe each molecule in a single ice with infrared spectroscopy. For this reason, much of the new work we report here relies heavily on the use of isotopically labeled compounds to separate the most-intense CO and ketene features. The great majority of our experiments involved the use of $^{13}$C- or $^{18}$O-enriched reagents, although a small number of deuterated compounds also were studied. The latter proved to give IR spectra with ketene features that overlapped with bands of other products, and so no experiments using deuteration-enriched reagents are reported here.

Table 2 summarizes some literature results for ketene’s IR spectrum. Absolute intensities have been measured for only a few features (Berg & Ewing 1991), with the $v_2$ band being by far the strongest, and comparable to the intense OCO asymmetric vibration of CO$_2$. Calculations for all of ketene’s mid-IR features show that the $v_2$ absorption is 5–10 times stronger than any other (Allen & Schaefer 1986). In our own work, this IR feature was found to be a sensitive probe of reaction chemistry in ices.

4.1. Ketene Formation in Apolar Ices

Previous experiments (Haller & Pimentel 1962) showed that the far-UV photolysis of N$_2$O at 20 K produces oxygen atoms that react with C$_2$H$_2$ to make ketene. We have extended and modified this experiment using apolar ices, and Figure 2 shows the results from three different samples. Spectrum (a) is after photolysis of O$_2$ + C$_2$H$_2$ with each element in its natural abundance. The lone IR peak near 2136 cm$^{-1}$ is at the position of both CO and ketene. Trace (b) is the spectrum produced by photolysis of an ice composed of $^{18}$O$_2$ + C$_2$H$_2$ and shows two well-resolved features. The one on the left is from labeled ketene ($H_2$CC$^{18}$O), while the one on the right is from C$^{18}$O. Spectrum (c) is from a photolyzed O$_2$ + $^{13}$C$_2$H$_2$ ice, with the assignments being $^{13}$CO for the leftmost peak and H$_2$C$^{13}$C for the rightmost.

Taken together, these experiments demonstrate that oxygen atoms, produced by far-UV decomposition of O$_2$, react with acetylene to make ketene. To support our spectral assignments, Table 3 gives the positions we measured for the strongest IR features of CO and ketene, along with isotopic shifts calculated by density functional theory at the B3LYP/6–31+G* level. Some of the ketene isotopologues we studied have been examined in...
Figure 2. Spectra of the UV photolysis (90 minutes) at \(\sim 15\) K of \(\text{O}_2 + \text{C}_2\text{H}_2\) ices, showing formation of ketene (\(\text{H}_2\text{C}=\text{C}=\text{O}\)) and \(\text{CO}\). The ices initially were composed of (a) \(\text{O}_2 + \text{C}_2\text{H}_2\), (b) \(^{18}\text{O}_2 + \text{C}_2\text{H}_2\), and (c) \(\text{O}_2 + ^{13}\text{C}_2\text{H}_2\). No IR features were present in this region before photolysis of the samples. In each case, the initial ice composition was \(\text{O}_2: \text{C}_2\text{H}_2 \approx 1:1\).

Table 2

<table>
<thead>
<tr>
<th>Vibrational Mode</th>
<th>Position(^a) (cm(^{-1}))</th>
<th>Intensity(^b) (cm molecule(^{-1}))</th>
<th>Approximate Description(^c)</th>
</tr>
</thead>
<tbody>
<tr>
<td>(\nu_1)</td>
<td>3043</td>
<td>...</td>
<td>C–H stretch</td>
</tr>
<tr>
<td>(\nu_2)</td>
<td>2133</td>
<td>(1.2 \times 10^{-16})</td>
<td>C=O stretch</td>
</tr>
<tr>
<td>(\nu_3)</td>
<td>1374</td>
<td>(2.7 \times 10^{-16})</td>
<td>H–C–H bend</td>
</tr>
<tr>
<td>(\nu_4)</td>
<td>1131</td>
<td>(4.6 \times 10^{-18})</td>
<td>C=C stretch</td>
</tr>
<tr>
<td>(\nu_5)</td>
<td>3140</td>
<td>...</td>
<td>C–H stretch</td>
</tr>
<tr>
<td>(\nu_6)</td>
<td>971</td>
<td>...</td>
<td>CH(_2) rock</td>
</tr>
<tr>
<td>(\nu_7)</td>
<td>440, 435</td>
<td>...</td>
<td>C=C=O bend</td>
</tr>
<tr>
<td>(\nu_8)</td>
<td>616</td>
<td>...</td>
<td>CH(_2) wag</td>
</tr>
<tr>
<td>(\nu_9)</td>
<td>529, 512</td>
<td>...</td>
<td>C=C=O bend</td>
</tr>
</tbody>
</table>

Notes.
\(^a\) Moore & Pimentel (1963).
\(^c\) Allen & Schaefer (1986).

Table 3

<table>
<thead>
<tr>
<th>Molecule</th>
<th>Observed Position (cm(^{-1}))</th>
<th>Observed Isotopic Shift(^b) (cm(^{-1}))</th>
<th>Calculated Isotopic Shift(^c) (cm(^{-1}))</th>
</tr>
</thead>
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<td>(^{12}\text{C})=(^{15}\text{C})=(^{16}\text{O})</td>
<td>2136</td>
<td>...</td>
<td>...</td>
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<tr>
<td>(^{12}\text{C})=(^{12}\text{C})=(^{16}\text{O})</td>
<td>2107</td>
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<td>28</td>
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<tr>
<td>(^{12}\text{C})=(^{13}\text{C})=(^{16}\text{O})</td>
<td>2071</td>
<td>65</td>
<td>65</td>
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<tr>
<td>(^{12}\text{C})=(^{12}\text{C})=(^{16}\text{O})</td>
<td>2109</td>
<td>27</td>
<td>26</td>
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</tr>
<tr>
<td>(^{12}\text{C}^{18}\text{O})</td>
<td>2088</td>
<td>48</td>
<td>53</td>
</tr>
<tr>
<td>(^{13}\text{C}^{18}\text{O})</td>
<td>2091</td>
<td>45</td>
<td>49</td>
</tr>
</tbody>
</table>

Notes.
\(^a\) Positions observed in our work for each molecule’s strongest IR feature.
\(^b\) Shifts are with reference to the positions of the \(^{12}\text{C}, ^{1}\text{H}, \text{and} ^{16}\text{O}\) isotopologues of ketene and CO.
\(^c\) Our own density functional calculations.

the gas or liquid phases in other laboratories (e.g., Hochstrasser & Wirz 1990), and in all cases our measured isotopic shifts are in excellent agreement with published laboratory results and our own calculated values. The results of Figure 2 served to determine spectral positions and acted as preliminary work for our studies of ion-irradiated ices.

Figure 3 shows IR spectra from the experiment represented by Figure 2(b), namely an \(^{18}\text{O}_2 + \text{C}_2\text{H}_2\) (1:1) ice at 15 K after several
Figure 3. Spectra showing $^{18}$O$_2$ + C$_2$H$_2$ (1:1) at $\sim$15 K after various stages of UV photolysis (bottom to top).

Figure 4. Spectra showing the similarity of reaction products from UV photolysis and proton irradiation, and from O$_2$ and CO$_2$ as sources of oxygen: (a) O$_2$ + $^{13}$C$_2$H$_2$ after 5 minute UV photolysis, (b) O$_2$ + $^{13}$C$_2$H$_2$ after $2 \times 10^{13}$ $p^*$ cm$^{-2}$, and (c) CO$_2$ + $^{13}$C$_2$H$_2$ after $1 \times 10^{14}$ $p^*$ cm$^{-2}$, with each ice initially being a 1:1 mixture.

far-UV exposures. The first band to appear was that of $^{18}$O-labeled ketene, on the left, followed by the fundamental band of C$^{18}$O. This observation is consistent with our hypothesis that the CO we found earlier (Moore & Hudson 1998) in irradiated acetylene-containing ices was from the sequence C$_2$H$_2$ $\rightarrow$ H$_2$CCO $\rightarrow$ CO. Not shown in the figure, but consistent with this sequence, is the observation that the ketene feature went through a rise and then a fall in intensity with continued photolysis, while the CO feature continued to grow (vide infra).

Figure 4 shows an important link between the preliminary photolysis work just described and our radiation studies. Traces (a) and (b) of Figure 4 demonstrate that the same IR features are found in UV-photolyzed and ion-irradiated ices, while trace (c) shows that irradiated CO$_2$, like O$_2$, produces oxygen atoms that react with acetylene to make ketene. These spectra unambiguously show that ketene formation occurs in both irradiated and photolyzed apolar ices.

4.2. Ketene Formation in H$_2$O-rich Ices

As preliminary work for our radiation experiments on H$_2$O-rich ices, we UV-photolyzed a sample made from H$_2$O + $^{13}$C$_2$H$_2$ ($\sim$5:1). The spectra that resulted are shown in Figure 5. Just as in Figure 3, ketene seems to form first, followed by CO. Figure 5 shows IR spectra from a proton irradiation of H$_2$O + $^{13}$C$_2$H$_2$ (5:1) at 18 K. The band growing in at 2090 cm$^{-1}$ is due to $^{13}$CO, while the adjacent feature is from H$_2^{15}$C$^{13}$O. That ketene again appeared first, followed by CO, seems clear. Also, this same figure shows that ketene’s IR feature rose and fell with dose, while the CO band continued to grow, supporting the assumed C$_2$H$_2$ $\rightarrow$ H$_2$CCO $\rightarrow$ CO sequence. Slowly warming this same irradiated ice (Figure 6) from 18 K to 165 K caused the ketene absorbance to decrease by about 75%, and essentially removed all of the $^{13}$CO absorption. Extending these observations, the reactions in irradiated H$_2$O + $^{12}$C$_2$H$_2$ ices studied earlier in our
4.3. Other Results

During the course of our work a few other observations were made that we wish to report. First, Bennett et al. (2005b) have published an extensive study of electron-irradiated CO$_2$ + C$_2$H$_4$ ice at 10 K, but did not report ketene formation. Their non-observation of ketene is probably due to their use of reagents having C, H, and O in their natural abundances. Figure 7 shows our own IR spectra from radiation experiments with $^{18}$O-enriched reagents, which confirm ketene formation in ices containing C$_2$H$_4$ (ethylene). These experiments emphasize the value of isotopically labeled materials in similar studies of ketene in astronomical ice analogs.

Related to this point, in both photolyzed and irradiated O$_2$ + C$_2$H$_2$ ices we observed not only CO and ketene synthesis, but also the production of CO$_2$. Isotopically labeled reagents were important in discriminating reaction chemistry from possible contamination by atmospheric CO$_2$.

Another irradiated-ice mixture briefly examined was solid-phase CO + CH$_4$. The experiments of Bennett et al. (2005a) did not include a detection of ketene, which we again suspect was due to the use of reagents with C, H, and O in natural abundance. In contrast, when we irradiated $^{13}$CO + $^{13}$CH$_4$ ices, $^{13}$C-labeled ketene was indeed observed. Besides ketene, we also found the ketenyl radical (HCCO) in several different experiments. Table 4 summarizes the isotopic shifts we observed for CO$_2$ and HCCO in our work.
Finally, to help elucidate the reaction paths leading to ketene we performed two types of scavenging experiments. Adding \(~1\%\) NH\(_3\), an H\(^+\) scavenger, to H\(_2\)O + C\(_2\)H\(_2\) ices before irradiation scarcely changed the experimental results. Similarly, adding \(~1\%\) N\(_2\)O, an electron scavenger, to H\(_2\)O + C\(_2\)H\(_2\) did not alter the final products. These experiments suggested that the observed products arose primarily through free-radical pathways, as opposed to acid–base chemistry (H\(^+\) transfers) or redox processes (\(e^-\) transfers). See Hudson & Moore (2000, and references therein) for another application of chemical scavengers to interstellar ice analogs.

5. DISCUSSION

5.1. Ketene Spectra and Chemistry

In general, the positions of IR absorptions in a mixed molecular solid are seldom the same as the positions of the individual components in their pure states (e.g., Knez et al. 2012; Palumbo et al. 1995). This led us to rely heavily on ketene’s isotopic shifts, as given in Table 3, in interpreting our results, with the good agreement between observed and calculated shifts supporting our ketene spectral assignments. A second line of support came from comparisons to the similar chemical systems in Table 5. The table gives products of triply bonded molecules after reaction with photochemically produced oxygen atoms, and in each case the resulting product is from an O-atom capture. Therefore, we took reactions (1) and (2) below as the source of ketene from O\(_2\) dissociation:

\[
\begin{align*}
\text{O}_2 & \rightarrow \text{O} + \text{O} \quad (1) \\
\text{O} + \text{C}_2\text{H}_2 & \rightarrow \text{H}_2\text{C}=\text{C}=\text{O} \quad (2)
\end{align*}
\]

Table 4:

<table>
<thead>
<tr>
<th>Initial Ice</th>
<th>Observed Position and Assignment (cm(^{-1}))</th>
<th>Observed Isotopic Shift(^b) (cm(^{-1}))</th>
<th>Literature Isotopic Shift(^b) (cm(^{-1}))</th>
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</thead>
<tbody>
<tr>
<td>O(_2) + C(_2)H(_2)</td>
<td>2341, CO(_2)</td>
<td>...</td>
<td>...</td>
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<tr>
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<tr>
<td>18O(_2) + C(_2)H(_2)</td>
<td>2307, C(_{18})O(_2)</td>
<td>34</td>
<td>35(^c)</td>
</tr>
<tr>
<td>O(_2) + 13C(_2)H(_2)</td>
<td>2275, 13CO(_2)</td>
<td>66</td>
<td>66(^c)</td>
</tr>
<tr>
<td>CO + CH(_4)</td>
<td>2026, HCCO</td>
<td>...</td>
<td>...</td>
</tr>
<tr>
<td>CO + CD(_4)</td>
<td>1997, DCCO</td>
<td>29</td>
<td>29(^d)</td>
</tr>
<tr>
<td>CO + 13CO + 13CH(_4)</td>
<td>2020, H(_{13})CCO</td>
<td>6</td>
<td>...(^d)</td>
</tr>
<tr>
<td>CO + CD(_4)</td>
<td>1968, H(<em>{13})C(</em>{13})CO</td>
<td>58</td>
<td>58(^d)</td>
</tr>
</tbody>
</table>

Notes.

\(^a\) Positions observed in our work for each molecule’s strongest IR feature.
\(^b\) Shifts are with reference to the positions of the \(^{12}\)C, \(^{1}\)H, and \(^{16}\)O isotopologues.
\(^c\) Vasserot et al. (2003) in neon matrices.
\(^d\) Forney et al. (1995) in neon matrices. No data for H\(_{13}\)CCO reported, but our own density-functional calculations gave isotopic shifts of 28, 5, and 62 cm\(^{-1}\) for DCCO, H\(_{13}\)CCO, and H\(_{13}\)C\(_{13}\)CO, respectively.

Table 5:

<table>
<thead>
<tr>
<th>Reaction</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>H-(\equiv)C-H + O \rightarrow H(_2)C=(\equiv)C=O</td>
<td>Haller &amp; Pimentel (1962)</td>
</tr>
<tr>
<td>H-(\equiv)N: + O \rightarrow HN=(\equiv)O</td>
<td>Crowley &amp; Sodeau (1989)</td>
</tr>
<tr>
<td>:(\equiv)O: + O \rightarrow O=(\equiv)O</td>
<td>Elsila et al. (1997)</td>
</tr>
<tr>
<td>:(\equiv):N: + O \rightarrow N=(\equiv):N</td>
<td>Elsila et al. (1997)</td>
</tr>
</tbody>
</table>
Analogous reactions between either NH or CH₂ (isoelectronic to O) and the four triply bonded molecules of Table 5 can be written (e.g., Milligan & Jacox 1963; Moore & Pimentel 1964).

Our CO₂ + C₂H₂ ice (Figure 4) also gave ketene, presumably in a process in which reaction (1) was replaced by CO₂ → CO + O, followed by reaction (2). We estimate that in these experiments, the maximum yield of ketene is ∼0.2% with respect to the initial abundance of C₂H₂. Our O₂ + C₂H₂ experiments gave a similar yield, 0.5%. These numbers will increase if one considers that ketene is also likely the source for the ¹³CO formed in these experiments. In the three ices containing ethylene (Figure 7), the oxidation sequence C₂H₄ → C₂H₂ was the source of acetylene for reaction (2).

Turning to our H₂O-rich samples (polar ices), although the ketene seen in ion-irradiated H₂O + C₂H₂ solids also may come from O-atom reactions, we tend to favor a different pathway. Radiolysis dissociates O₂ into O atoms, hydroxyl radicals (OH), and a number of other reactive species (e.g., H₂O*, e⁻, HO₂). Acetylene undergoes H and OH addition reactions (Hudson & Moore 1997) to make ethylene (C₂H₄), ethane (C₂H₆), and ethanol (C₂H₅OH) through

\[ \text{C}_2\text{H}_2 \rightarrow \text{C}_2\text{H}_4 \rightarrow \text{C}_2\text{H}_6 \] (3)

but vinyl alcohol, C₂H₅OH, forms if the initial reaction in reaction (3) or (4) involves H and OH radicals. Earlier we identified vinyl alcohol as a product in irradiated H₂O + C₂H₂ ices, and observed that its abundance first rose and then fell with increasing radiation dose (Hudson & Moore 2003). This fall in vinyl alcohol’s abundance has multiple causes, including the molecule’s isomerization into acetaldehyde (Erlenmeyer 1881). An additional destructive path is vinyl alcohol’s direct decomposition into ketene by H₂ elimination, the assumed sequence being as shown in reaction (5).

\[ \text{C}_2\text{H}_3(\text{OH}) \rightarrow \text{C}_2\text{H}_2=\text{C}=:\text{O} + \text{H}_2. \] (5)

An alternate route to ketene from H₂O + C₂H₂ ices can be envisioned through HC(=O)CH₃ (acetaldehyde) in the sequence

\[ \text{HC}≡\text{CH} \rightarrow \text{HC}≡\text{C}=\text{CH(=O)} \rightarrow \text{HC}=\text{C}=\text{O} + \text{H}_2 \]

although we have not yet investigated this pathway in detail. Similar reactions will apply to UV-photozed H₂O + C₂H₂ ices. See Hawkins & Andrews (1983) for a discussion of photoelimination of H₂ from ketene in inert-gas matrices.

As already stated, reaction (5) competes with vinyl alcohol destruction along other paths. Such competing reactions, and our earlier estimate (Hudson & Moore 2003) of ∼40% for the conversion of acetylene into vinyl alcohol in an H₂O-rich ice, suggest that any yield of ketene must be less than this value. For the radiation experiment of Figure 6, half of the C₂H₂ initially present was lost after a fluence of about 4 × 10¹⁴ p⁺ cm⁻², which also corresponded with roughly the maximum amount of ketene observed. Comparing the band area of our ketene produced in H₂O-rich ices to band areas in ices of similar thickness, we estimate a maximum acetylene-to-ketene yield of ∼0.5% for the experiment of Figure 6, which is comparable to the values obtained in our CO₂ + C₂H₂ and O₂ + C₂H₂ mixtures. This yield will at least double if we assume that ketene is also the source for all of the ¹³CO formed in this same experiment (IR band shown in Figure 6). Note that nearly all such calculations of reaction yields in ices are hindered by the lack of precise measurements for such quantities as IR band strengths, densities, and refractive indices (Satorre et al. 2008).

The focus of our work has been ketene’s formation in interstellar ices, but we also have observed several ketene decomposition products such as CO, CO₂, and ketenyl (HCCO). Ketene dissociation can account for CO, which in reaction with either O atoms or OH radicals will give CO₂. Direct photo- or radiolytic dissociation of ketene,

\[ \text{H}_2\text{C}≡\text{C}=\text{O} \rightarrow \text{HCCO} + \text{H} \] (7)
or H-atom abstraction

\[ \text{H} + \text{H}_2\text{C}≡\text{C}=\text{O} \rightarrow \text{H}_2 + \text{HCCO} \] (8)

will generate ketenyl radicals. We note in passing that the ketenyl radical (H–C≡C=O) and what is called the ethynylxy radical (H=C≡C=O) are related by way of resonance (H–C≡C=O ↔ H=C≡C=O), and so are the same species. We also caution that what modern chemists call ketenyl occasionally has been termed “ketyl” in the astronomical literature (e.g., Turner & Sears 1989; Ruiterkamp et al. 2007).

5.2. Ketene and Astrochemistry

The radiation doses that icy grain mantles in dense interstellar clouds are expected to receive are on the order of 3 eV molecule⁻¹ over a million years (Moore et al. 2001). The spectra of our Figure 6 are for an H₂O + C₂H₂ ice with the conversion between fluence and dose being 1 × 10¹⁴ p⁺ cm⁻² ≈ 1 eV molecule⁻¹. Therefore, the fact that we see ketene formation after ∼1 eV molecule⁻¹ suggests that ketene should readily form in dark clouds from reactions between H₂O-ice and trapped C₂H₂. A similar conclusion follows for CO₂-rich ices containing C₂H₂, based on Figure 4. For photolysis experiments, the time to convert acetylene to ketene can be estimated from Ehrenfreund et al. (2001), who determined that 32 million years of UV exposure in a dark cloud was comparable to about 1000 s of photolysis with their lamp. For our own UV experiments, conducted at a lower photon flux, the conversion is about 3000 s for 32 million years in a dense molecular cloud. Since our ice experiments readily showed ketene formation after ∼10 minutes the equivalent reaction time in a cold, dense interstellar cloud would be around 6 million years.

Taken together, our experiments demonstrate the relative ease of radiation synthesis of ketene in interstellar environments, and in both polar and apolar ices. For polar ices, only H₂O, C₂H₂, and a suitable source of energy, are needed; for apolar ices, O atoms and C₂H₂ suffice. Acetylene, water, oxygen atoms, and cosmic, or UV, radiation are all part of the interstellar environment and so ketene is expected in interstellar ices. Moreover, ketene has long been suspected to form by the decomposition of most any molecule with the −CH₂=C(=O)− group (Wilsmore 1907), and so the gas-phase presence of many known interstellar molecules necessarily suggests the presence of ketene as well. Among these molecules are acetic acid, acetone, acetaldehyde, and acetamide, all shown in Figure 8. Conversely, reactions that are the reverse of those in Figure 8 might prove to be a low-temperature source of acetic acid and the other interstellar molecules shown.

A possible complication arising from our work on ketene formation is that the overlap of the IR bands of solid CO and solid ketene near 2140 cm⁻¹ (4.67 μm) could require modifications...
to the reported solid-phase CO abundances in interstellar ices (Gibb et al. 2004). However, we believe that such changes are likely to be small. If the cometary abundance of C₂H₂, \( \sim 0.3\% \) of H₂O (Mumma & Charnley 2011), is comparable to the abundance of acetylene in an interstellar ice then even complete acetylene-to-ketene conversion (\( H-C≡C-H \rightarrow H₂C≡C≡O \)) would not significantly alter the reported \( >3\% \) CO seen in interstellar ices. A possible exception might be interstellar clouds that have warmed sufficiently to remove some solid CO from grains by sublimation, leaving ketene as a more-significant contributor to the IR band near 2140 cm\(^{-1}\) (4.67 \( \mu \)m).

While much of our work has focused on interstellar chemistry, applications to solar system astrochemistry are possible. A radiation dose of about 5 eV molecule\(^{-1}\) is expected at a 5 m depth to an Oort-Cloud comet over the age of the solar system, \( \sim 4.6 \times 10^9 \) yr (Moore et al. 2001). This means that even if ketene were not to form in interstellar ices, its post-accretion synthesis from cometary H₂O-ice and C₂H₂ trapped in a comet’s nucleus is expected. Warming of a cometary ice will release ketene, primarily at the temperature where the H₂O-ice undergoes sublimation (Hudson & Donn 1991). The extent to which any such cometary H₂C≡C≡O would subsequently dissociate in the gas phase, contributing to a CO distributed source, would depend on the available UV light, a possibility that remains to be explored. This same pathway for ketene formation also may have implications for TNO chemistry since both CH₄ and CO have been seen as ices on Pluto (Grundy & Buie 2001).

A final application of our work concerns meteoritic chemistry. Cooper et al. (2011) have analyzed the organic component of a carbonaceous meteorite (Murchison) and identified pyruvic acid, citric acid, and several other metabolism precursors. The interstellar synthetic paths for these materials remain unknown, but an intriguing possibility is that they all originate with the formation of ketene and its subsequent reaction to make pyruvonitrile, H₂C≡C(=O)–CN. The decomposition of pyruvonitrile to make ketene has been demonstrated at \( \sim 10 \) K in argon matrices (Guennoun et al. 2005). Now that we have shown that ketene is easily formed in ices from interstellar molecules, new work is needed to characterize the reverse process, ketene \( \rightarrow \) pyruvonitrile. In short, the laboratory results reported in the present paper demonstrate ketene formation in ices and thereby lend support to the possible formation of far-more complex organics as outlined elsewhere (Cooper et al. 2011). Details of the low-temperature “ketene \( \rightarrow \) pyruvonitrile \( \rightarrow \) complex organics” process remain to be studied.

6. SUMMARY AND CONCLUSIONS

Our experiments have shown that ketene formation can occur in polar and apolar ices dominated by known interstellar molecules such as C₂H₂, CH₄, H₂O, CO₂, and O₂. Both radiolysis and photolysis can initiate ketene syntheses, and with radiolytic yields on the order of 0.1%–1%. Although multiple reaction pathways to ketene can be envisioned, our experiments suggest that a free-radical reaction mechanism is operative, implying that ketene formation will not be blocked by solid-phase NH₃ (a base) in interstellar ices. In polar ices, ketene formation is an intermediate step in the solid-phase oxidation of C₂H₂ to CO, the sequence being

\[
C₂H₂ \rightarrow CH₂CH(OH) \rightarrow H₂C≡C≡O \rightarrow CO.
\]

In apolar (H₂O-poor) ices, both CO₂ and O₂ can serve as the source of O atoms for ketene formation and both C₂H₂ and C₂H₄ can serve as the carbon source. The thermal stability of ketene in H₂O-rich ices is such that some of it will be retained until the H₂O-ice sublimes. A similar conclusion applies to ketene in CO₂-rich ices.

To extend these experiments, better measurements of reaction yields are needed, which in some cases will mean new work to reduce uncertainties in the relevant IR-band strengths. The reactions of Figures 1 and 8 also need further study in the solid state, as does ketene formation in more-complex ice mixtures. In all cases, IR spectroscopy experiments should use isotopically enriched reagents to distinguish CO from ketene. Experiments that do not, and predictions from computational models arising from them, should be viewed with caution.

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