Tropical Atlantic dust and smoke aerosol variations related to the Madden-Julian Oscillation in MODIS and MISR observations

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[1] In this study, Moderate Resolution Imaging Spectroradiometer (MODIS) fine mode fraction and Multi-angle Imaging SpectroRadiometer (MISR) nonspherical fraction data are used to derive dust and smoke aerosol optical thickness (τdust and τsmoke) over the tropical Atlantic in a complementary way: due to its wider swath, MODIS has 3–4 times greater sampling than MISR, but MISR dust discrimination is based on particle shape retrievals, whereas an empirical scheme is used for MODIS. MODIS and MISR show very similar dust and smoke winter climatologies. τdust is the dominant aerosol component over the tropical Atlantic, accounting for 40–70% of the total aerosol optical thickness (AOT), whereas τsmoke is significantly smaller than τdust. The consistency and high correlation between these climatologies and their daily variations lends confidence to their use for investigating the relative dust and smoke contributions to the total AOT variation associated with the Madden-Julian Oscillation (MJO). The temporal evolution and spatial patterns of the τdust anomalies associated with the MJO are consistent between MODIS and MISR: the magnitude of MJO-related τdust anomalies is comparable to or even larger than that of the total τ, while the τsmoke anomaly represents about 15% compared to the total, which is quite different from their relative magnitudes to the total τ on the climatological time scale. This suggests that dust and smoke are not influenced by the MJO in the same way. Based on correlation analysis, dust is strongly influenced by the MJO-modulated trade wind and precipitation anomalies, and can last as long as one MJO phase, whereas smoke is less affected.


1. Introduction

[2] The Madden-Julian Oscillation (MJO) [Madden and Julian, 1971, 1972] is the dominant form of intraseasonal (30–90 day) variability in the tropical atmosphere. It is characterized by slow (~5 m s−1) eastward-propagating, large-scale oscillations in the tropical deep convection over the equatorial Indian Ocean and western Pacific during boreal winter (November–April) [Lau and Waliser, 2005; Zhang, 2005]. Recently, there is strong emerging interest in the impacts of the MJO on atmospheric composition [Tian and Waliser, 2011], such as aerosol [Tian et al., 2008, 2011; Reid et al., 2012], ozone [Tian et al., 2007; Weare, 2010; Li et al., 2011], carbon dioxide [Li et al., 2010], and carbon monoxide [Wong and Deessler, 2007].

[3] Tian et al. [2008] first examined the aerosol variability related to the MJO using global aerosol products from multiple sensors on various satellite platforms. That study revealed large intraseasonal variations in the satellite-derived aerosol products over the tropics, though the interpretation in terms of actual aerosol behavior was ambiguous as inconsistent aerosol vs. MJO-wet-phase relationships are found between different satellite aerosol data sets. Possible reasons leading to the inconsistency include the aerosol humidification effect, wet deposition, low-level wind effect, biological production, sampling effect, and cloud contamination. Tian et al. [2011] further investigated the MJO-related aerosol variability over the tropical Atlantic Ocean using the aerosol optical thickness (AOT) product from the MODIS (Moderate Resolution Imaging Spectroradiometer) on the Aqua satellite. They suggested that the MJO-related intraseasonal variance accounts for about 25% of the total AOT variance over the tropical Atlantic. They also found that the MJO modulates the Atlantic aerosol variation primarily through its influence on the Atlantic low-level zonal winds. Given the potential predictability of the MJO extending to 2–4 weeks [e.g., Waliser, 2005], the study by Tian et al. [2011] implies that the Atlantic aerosol concentration may be predictable with...
lead times of 2–4 weeks, which in turn may lend important
guidance to predicting air quality, dust storm activity, and
ocean nutrient deposition over the Atlantic Ocean.

[5] Nevertheless, Tian et al. [2011] examined only the
total AOT anomalies and did not consider the contribution
of different aerosol types to the total AOT anomalies. It is
well known that the aerosol over the tropical Atlantic Ocean
in the boreal winter season is a mixture of mineral dust from
the Sahara desert and the Sahel region, biomass burning
smoke from the Sahel and African savanna regions, and
marine aerosol (primarily sea salt and secondly sulfate aerosols)
from the ocean surface [Kaufman et al., 2002, 2005a, 2005b].
Because dust, biomass burning smoke and marine aerosols
play very different roles in the radiative forcing and cloud
formation process, it is of great interest to partition the total
aerosol into individual aerosol components and examine the
MJO-related variability in each aerosol type.

[5] In this study, we use two independent satellite data sets,
MODIS and MISR (Multi-angle Imaging Spectroradiometer),
to investigate the intraseasonal variability of individual aerosol
components related to the MJO over the tropical Atlantic.
These two data sets both have their advantages and disadvan-
tages, thus are complementary to each other: MODIS has more
frequent sampling than MISR due to its wider swath, but the
scheme used to distinguish dust and smoke is subject to some
major assumptions. MISR discriminates dust aerosol using
actual particle shape information, however, it cannot distin-
guish smoke from other spherical components, and it has only
1/4 to 1/3 times sampling of MODIS.

[5] The rest of this paper is organized as follows. Section
2 describes the MODIS and MISR data, and the methodolo-
y used to derive specific aerosol components from the
total AOT. The climatology of specific aerosol components
as well as the comparison of dust aerosol between MODIS
and MISR is presented in section 3 to examine the fidelity
of the methods described in section 2. The main results of
this paper, the MJO-related dust and smoke aerosol anom-
ali es and correlations with MJO-related wind and precipita-
tion anomalies, are presented in section 4. Conclusions and
discussions are presented in section 5.

2. Methodology and Data Description

2.1. Moderate Resolution Imaging Spectroradiometer

have suggested that satellite data distinguishing fine-mode
aerosols from coarse-mode aerosols could be used to sepa-
rate the aerosol into specific types, because different aerosol
types (e.g., smoke, dust, and marine aerosols) have different
fine mode fraction (f) values. f is the fraction of total
midvisible AOT contributed by the fine-mode aerosols. In
MODIS, the fine-mode aerosols refer to the aerosols with a
size distribution of radii centered between 0.1 and 0.25 μm,
whereas the coarse-mode aerosols have radii centered
between 1 and 2.5 μm. In their series of studies [Kaufman
et al., 2002, 2005a, 2005b], they developed empirical for-
mulae to estimate dust and smoke aerosols using MODIS/
Terra AOT and f observations. These formulae have since
been widely adopted by the community to understand aerosol
types and their climate forcing [e.g., Bellouin et al., 2005; Yu
et al., 2009, among many others]. With the two constraints that
both the total AOT and its fine mode fraction can be
partitioned into contributions from the dust, smoke, and
marine aerosol components, we have the following equations:

\[ \tau = \tau_{\text{dust}} + \tau_{\text{smoke}} + \tau_{\text{mar}} \] (1)

\[ f \times \tau = f_{\text{dust}} \times \tau_{\text{dust}} + f_{\text{smoke}} \times \tau_{\text{smoke}} + f_{\text{mar}} \times \tau_{\text{mar}} \] (2)

[8] Rewriting (1) and (2), we get

\[ \tau_{\text{dust}} = \frac{[\tau \times (f_{\text{smoke}} - f) - \tau_{\text{smoke}} \times (f_{\text{smoke}} - f_{\text{mar}})]}{(f_{\text{smoke}} - f_{\text{dust}})} \] (3)

\[ \tau_{\text{smoke}} = \frac{[\tau \times (f - f_{\text{dust}}) - \tau_{\text{dust}} \times (f_{\text{mar}} - f_{\text{dust}})]}{(f_{\text{smoke}} - f_{\text{dust}})} \] (4)

[9] In these equations, \( \tau \) denotes total midvisible AOT
(τ and AOT are used interchangeably in this paper), and the
subscripts “dust”, “smoke”, and “mar” indicate dust, smoke,
and marine aerosol components, respectively. With \( \tau \) and \( f \) being
MODIS measurements, \( \tau_{\text{dust}} \) and \( \tau_{\text{smoke}} \) can be computed
directly if \( f_{\text{dust}}, f_{\text{smoke}}, f_{\text{mar}} \) are known.

[10] In the Kaufman et al. study, \( f_{\text{dust}}, f_{\text{smoke}}, \) and \( f_{\text{mar}} \) are
assumed to be constant, and were derived by averaging the
MODIS/Terra f observations over selected regions and time
periods where one specific aerosol type dominates, with
their uncertainties estimated from these selected measure-
ments: \( f_{\text{dust}} = 0.5 \pm 0.05, f_{\text{smoke}} = 0.9 \pm 0.05, \) and \( f_{\text{mar}} = 0.3 \pm 0.1 \) [Kaufman et al., 2005a]. However, the actual \( f \) values
vary with season and location [e.g., Maring et al., 2003; Yu
et al., 2009], thus large uncertainties are expected from
using constant \( f_{\text{dust}}, f_{\text{smoke}}, f_{\text{mar}} \) when applying Kaufman’s
formula. More discussion of this issue as well as the sensi-
tivity of our results to the \( f \) values adopted will be given in
section 4.3.

[11] The marine AOT, \( \tau_{\text{mar}} \), depends strongly on surface
wind speed, as its primary component is sea-spray salt
[Huang et al., 2010; Smirnov et al., 2003]; it is estimated
using the empirical formula in Kaufman et al. [2005a]

\[ \tau_{\text{mar}} = 0.007W + 0.02 \] (5)

[12] Here \( W \) is the surface (10 m) wind speed from the
European Centre for Medium-Range Weather Forecasts
(ECMWF) ERA-Interim reanalysis [Dee et al., 2011]. The
global mean value of \( \tau_{\text{mar}} \) is around 0.06 ± 0.005 [e.g.,
Kaufman et al., 2001].

[13] In this study, daily AOT and f measurements at
0.55 μm from the MODIS/Aqua Level-3 Collection 5.1 data
set [Remer et al., 2005; Levy et al., 2009] on 1° × 1° spatial
 grids are used. The uncertainties of \( \tau \) are ±(0.03 + 0.05τ)
over ocean and ±(0.05 + 0.15τ) over land. Over ocean, the
uncertainties of \( f \) are large for low AOT (\( \tau < 0.15 \)) but
typically less than about 20% for large AOT (\( \tau > 0.15 \))
[Kleidman et al., 2005; Remer et al., 2005]. Over land,
MODIS does not provide any quantitative information about
the aerosol size [Levy et al., 2010]. As a result, we use the \( \tau \) and \( f \) data
over the ocean only. The period of 4 July 2002 to 1 June 2009
is used for consistency with the study by Tian et al. [2011].

[14] Note that we use MODIS/Aqua aerosol data rather
than MODIS/Terra data even though MISR is on board
Terra. One reason is that we want to be consistent with the
study in Tian et al. [2011]. More importantly, MODIS/Terra
and MISR aerosol retrievals have very little spatial overlap due to the exclusion of Sun glint regions over the ocean for MODIS. As a result, there are far fewer days on which both MODIS/Terra and MISR made observations of a given region, compared to the situation between MODIS/Aqua and MISR. Because the direct comparison between daily MODIS and MISR aerosol data serves as an essential part of this study, we chose to use MODIS/Aqua data. Nevertheless, the difference between the results based on MODIS/Terra and those based on MODIS/Aqua is found to be negligible for our application.

[15] Data rejection has been performed on the MODIS $\tau$ and $f$ observations as discussed in Appendix A. It is noted that we applied equations (3) and (4) in a stricter way than what was originally done by Kaufman et al. [2005a, 2005b]. Nevertheless, it should be noted that although the data rejection we applied removes a large number of suspicious observations, doing so does not in itself guarantee the accuracy of the remaining points, given that $f_{\text{dust}}, f_{\text{smoke}},$ and $f_{\text{mar}}$ are assumed constant, and $\tau_{\text{mar}}$ is estimated from an empirical formula. For the former factor, sensitivity tests are performed to examine the extent to which our results are sensitive to the variations in $f$ values. For the latter, the deviations of the actual $\tau_{\text{mar}}$ (mainly sea salt) compared to that empirically computed will be interpreted primarily as changes in dust amount, thus affecting the accuracy of $\tau_{\text{dust}}$. However, as will be presented next, MISR can distinguish dust from sea salt based on retrieved particle shape, thus $\tau_{\text{dust}}$ derived from MISR is not subject to the ambiguity between dust and sea salt that affects Kaufman's method. In this sense, the examination of MISR aerosol observations provides validation for the more extensive MODIS data set, in addition to offering actual results independent of MODIS.

2.2. Multi-Angle Imaging SpectroRadiometer

[16] Viewing the Earth simultaneously at nine widely spaced angles, MISR/Terra provides global coverage every 7–9 days. The variations between the reflectance acquired from a range of observation angles can be interpreted (with appropriate models) in terms of aerosol properties such as particle size, shape, and single-scattering albedo [Kahn et al., 1998, 2001; Chen et al., 2008]. In particular, MISR’s sensitivity to the characteristics of the aerosol scattering phase function enables it to distinguish between the nonspherical and spherical particles, and thus provides a possible way to separate mineral dust aerosols from other aerosol components. A series of studies has explored the ability of MISR to retrieve mineral dust properties theoretically [Kahn et al., 1997, 2001; Kalashnikova et al., 2005; Kalashnikova and Kahn, 2006], the sensitivity of the theoretical results, as well as the application of nonspherical dust models for five Saharan dust field events over the Atlantic Ocean [Kalashnikova and Kahn, 2008].

[17] Assuming the mineral dust is all nonspherical and the nonspherical part of AOT is all from dust, MISR $\tau_{\text{dust}}$ can be directly computed as the nonspherical fraction of the total AOT:

$$\tau_{\text{dust}} = \tau \times \frac{\text{rac}}{\text{rac}_{\text{nonspherical}}} \quad (6)$$

[18] For MISR, smoke and marine aerosols (both sea salt and sulfate aerosols) contribute to the spherical part of the total AOT. (Refer to Table 3 in Kahn et al. [2001] for more details on shape categories of different aerosol components.) If $\tau_{\text{mar}}$ is again taken as computed from equation (5), $\tau_{\text{smoke}}$ can be approximated as the difference between the total $\tau$ and $\tau_{\text{dust}}$ and $\tau_{\text{mar}}$

$$\tau_{\text{smoke}} = \tau - \tau_{\text{dust}} - \tau_{\text{mar}} \quad (7)$$

[19] However, it should be kept in mind that unlike MISR $\tau_{\text{dust}},$ the MISR $\tau_{\text{smoke}}$ derived this way, and all the MODIS particle type distinctions, are not independent observations based on actual physical constraints, and their accuracy is subject to the large uncertainties due to the empirically calculated $\tau_{\text{mar}}$.

[20] We use MISR-derived Level 3 daily $\tau$ and nonspherical fraction at 0.558 $\mu$m on $0.5^\circ \times 0.5^\circ$ spatial grids during the same time period as the MODIS data. MISR $\tau$ and $\tau_{\text{dust}}$ are averaged onto $1^\circ \times 1^\circ$ grids to compare with MODIS results. A number of validation studies have shown that overall, about 70% to 75% of MISR AOT retrievals, their retrieval errors fall within 0.05 or 20% $\times$ AOT, and about 50% to 55% are within 0.03 or 10% $\times$ AOT, except at sites where mixed dust and smoke are commonly found [Kahn et al., 2010]. Particle property validation suggests that expected MISR sensitivity to the spherical versus nonspherical particles is about 20% for AOT above 0.15, and diminishes for midvisible AOT below this value [Kahn et al., 1997; Kalashnikova et al., 2005; Kalashnikova and Kahn, 2006]. Thus, we only use $\tau \geq 0.15$ to calculate $\tau_{\text{dust}}$ from equation (6). With this $\tau$ cutoff, it is found that more than 70% of the observations still remain for calculation over most of the tropical Atlantic. Note that generally, MISR $\tau$ retrievals are available about 15% of the time, except over the convectively active regions (decreased to about 10%). Although it has never been explicitly addressed, MODIS sensitivity to particle properties, e.g., $f,$ also diminishes at low AOT (implied in Figure 7 in Kahn et al. [2009]), and MISR actually has greater sensitivity at low AOT than MODIS due to the long atmospheric paths observed by its steeper-viewing cameras.

[21] Note that MISR also retrieves aerosol particle size information, thus conceptually it is possible to follow the same method utilized for MODIS to separate the total AOT into specific types in MISR. However MISR categorizes the aerosol particles into three bins: “fine” (particle radii $< 0.35 \mu$m), “medium” (radii between 0.35 and 0.7), and “large” (radii $> 0.7 \mu$m) modes, instead of two bins as “fine” vs. “coarse” in MODIS. Thus, applying Kaufman’s formula to MISR would require considerable additional work, but not necessarily lead to greater insight, because similar assumptions would be required to apply the MODIS size-discrimination method to MISR. Furthermore, the different radii range for MODIS and MISR “fine” mode would make it impossible to cross-validate the $f$ values between these two data sets. Most importantly, the retrieval of $\tau_{\text{dust}}$ in MISR is based on actual aerosol physical property, thus considered to be preferable to Kaufman’s method. Therefore, using MISR aerosol size information to separate different aerosol components following Kaufman’s method is not only beyond the scope of current paper, but also less desirable.
3. Comparison of Aerosol Components Between MODIS and MISR

In this section, the AOTs of individual aerosol components over the tropical Atlantic Ocean are examined, and the MODIS and MISR results are compared. Climatological maps are examined first, to investigate how well the methods presented in section 2 capture the basic features of the long-term mean.

3.1. Aerosol Winter Climatology

3.1.1. Moderate Resolution Imaging Spectroradiometer

The MODIS winter climatological mean (November to April during 2002–2009) aerosol maps over the tropical Atlantic Ocean are shown in Figure 1. The climatological mean features a zonally oriented, optically thick aerosol plume centered at around 5°N–8°N stretching across the equatorial Atlantic Ocean. The magnitude and latitudinal extent are greatest over the eastern equatorial Atlantic with the maximum up to 0.5 and gradually decrease westward toward the central and western equatorial Atlantic, as expected for an aerosol plume that originates in Africa (Figure 1a). Note that due to data rejection performed in this study, Figure 1a is highly similar but not identical to the color shadings in Figure 1b in Tian et al. [2011] although they display the same quantity: the magnitude of former is
about 80% of the latter over the equatorial region. Nevertheless, they have a spatial correlation of 0.98.

The spatial pattern of climatological $\tau_{dust}$ closely resembles the $\tau$ pattern, with maximum $\tau_{dust}$ (about 0.35) found along the west coast of Africa, and a gradual decrease toward the central and western equatorial Atlantic (Figure 1b). Over the equatorial Atlantic Ocean, where aerosol loading is high ($\tau > 0.15$), dust is the dominant aerosol component, contributing more than 50%, and as much as 75%, to $\tau$ (Figure 1c).

Compared to $\tau_{dust}$, $\tau_{smoke}$ is significantly weaker. Plumes of fine-mode-dominant aerosol are found originating from the African (biomass burning smoke) as well as the South American continents (smoke or air pollution) (Figure 1d). A contribution of more than 20% to total $\tau$ is found over the eastern tropical Atlantic, whereas the contribution is less than 20% over the western part (Figure 1e). Figure 1f shows that $\tau_{mar}$ is very small (about 0.04) over the Atlantic intertropical convergence zone and the west coast of Africa because of the weak trade winds. Over the clean subtropical Atlantic, marine aerosol is the dominant component (>50%) due to a lack of dust and smoke aerosols over these regions.

These results indicate that the major aerosol plume over the equatorial Atlantic Ocean in Figure 1a is the dust originating in the Sahara desert and the Sahel region, with some contribution from biomass burning smoke originating in the Sahel and African savanna. These aerosol distributions are generally consistent with previous observational results [e.g., Husar et al., 1997; Kaufman et al., 2005a; Huang et al., 2010], lending some confidence in estimating dust and smoke aerosols based on equations (3) and (4).

### 3.1.2. Multi-Angle Imaging SpectroRadiometer

Multi-angle Imaging SpectroRadiometer winter climatologies of $\tau$, $\tau_{dust}$, and $\tau_{smoke}$ are shown in Figure 2. Recall that for these results only cases with $\tau \geq 0.15$ are used due to the fact that sensitivity to shape diminishes when aerosol concentration is low. Different data rejection is applied to MODIS $\tau$, as discussed in the Appendix A, thus, we do not expect Figure 2a to be the same as Figure 1a. Nevertheless, we do find that they have highly consistent spatial patterns and comparable magnitudes. The climatological MISR $\tau_{dust}$ map (Figure 2b) is also very similar to the MODIS $\tau_{dust}$ (Figure 1b). The MISR $\tau$ and $\tau_{dust}$ are slightly larger than the MODIS counterparts partly due to MISR’s exclusion of low aerosol cases, but more important reasons will be addressed in section 3.2. It is seen that MISR $\tau_{dust}$ contributes more than 40%, up to more than 60%, to the total $\tau$ for cases with $\tau \geq 0.15$, further confirming that dust is the dominant aerosol component over the equatorial Atlantic Ocean (Figure 2c). It is also noted that MISR has 10% to 15% lower dust fraction over the equatorial Atlantic. Again, this difference could be partly attributed to the $\tau$ cutoff applied to MISR, but more importantly could result from
both uncertainties involved in the derivation of MODIS and MISR $\tau_{\text{dust}}$, especially in MODIS, given the assumptions used. We return to this point in section 3.2. The climatological MISR $\tau_{\text{smoke}}$ has a 10% larger contribution to total $\tau$ than MODIS $\tau_{\text{smoke}}$ over the equatorial Atlantic (Figures 2d and 2e compared to Figures 1d and 1e).

[30] The overall consistency between the MODIS and MISR aerosol climatology, representing two independent satellite data sets and methods used to derive the dust and smoke aerosols over the tropical Atlantic Ocean, gives us some confidence in applying the results. However, the consistency in climatology does not guarantee their consistency on shorter time scales, for example, on the intraseasonal time scale of importance here. Therefore, in the next subsection we directly compare the coincident daily MODIS and MISR $\tau_{\text{dust}}$ over the tropical Atlantic Ocean.

3.2. Comparison of Coincident MODIS and MISR Dust Aerosols

[30] The correlation between the daily MODIS and MISR $\tau_{\text{dust}}$ for all their coincident days (minimum of 5) during 2002–2009 is calculated over the tropical Atlantic. Here the term “coincident” simply means the MODIS and MISR observations fall within a same grid box and on a same day, which is less strict than the usual definition used by the satellite community. We use the entire years for 2002–2009 instead of winters only to obtain as many coincident days as possible.

[31] Overall, the MODIS and MISR $\tau_{\text{dust}}$ are well correlated in the regions of primary interest here, where dust dominates. The correlation is systematically higher in the north Atlantic region (typically around 0.7 or larger) than in the south Atlantic where the correlation is typically less than
0.5, with quite a few spots less than 0.2, or even negative (Figure 3a). Note that there are more coincident days over the north Atlantic (Figure 3b). Furthermore, the correlation is higher when it is closer to the African continent, and decreases gradually as the dust is transported away from the source region. These results suggest that MODIS and MISR dust agree with each other quite well over heavy dust load regions, whereas they are less consistent over the regions with less frequent dust occurrence or small dust aerosol concentration.

Figure 3a shows that overall, the MODIS and MISR \( \tau \) dust are highly correlated; however the correlation coefficient does not provide information on the \( \tau \) dust magnitude. Therefore, the time series of MODIS and MISR \( \tau \) dust as a function of their coincident days at a representative grid point (29.5°W, 19.5°N), as well as averaged within the 5° × 5° and 10° × 10° grid boxes centered at this point, are shown (Figures 3c–3e). Again it is found that the MODIS and MISR \( \tau \) dust are highly correlated on a daily basis, and it also reveals that \( \tau \) dust is systematically greater for MISR than MODIS.

Figure 4. (a) MISR and MODIS \( \tau \) difference averaged over coincident days during 2002–2009. (b) Same as Figure 4a, but for \( \tau \) dust. (c) Number of coincident MODIS and MISR observations (black line), averaged MODIS \( \tau \) (thick blue line), MISR \( \tau \) (thick red line), MODIS \( \tau \) dust (thin blue line), and MISR \( \tau \) dust (thin red line) over 20°S–30°N, 60°W–20°E during 2002–2009 as a function of binned MISR \( \tau \) (binned by every 0.01).

[32] Multi-angle Imaging SpectroRadiometer \( \tau \) dust is larger than MODIS \( \tau \) dust over most of the tropical Atlantic when averaged for coincident days (about 0.03 larger averaged over the basin), except in some regions north of the equator (Figure 4b). This difference can be traced to the difference between MISR and MODIS total AOT (Figure 4a). Over almost the entire tropical Atlantic, coincident MISR AOTs are systematically larger than the MODIS ones (up to 0.08, and about 0.04 when averaged for the Atlantic basin). This is consistent with previous studies in which MISR AOT is found to be generally larger than MODIS AOT over water [e.g., Abdou et al., 2005; Kahn et al., 2010]. Further examination of coincident MODIS and MISR \( \tau \) and \( \tau \) dust binned against the MISR \( \tau \) reveals that larger MISR \( \tau \) is found for the entire \( \tau \) spectrum, whereas MISR \( \tau \) dust is larger when the aerosol concentration is relatively low (Figure 4c). Note that lower aerosol concentrations are overwhelmingly more frequent (black line in Figure 4c), thus the averaged MISR \( \tau \) dust is larger than the averaged MODIS \( \tau \) dust. As seen above, the
The difference between MODIS and MISR can be traced to the difference between them. However, this is not the only reason, as, unlike MISR, MISR dust is not systematically larger than MODIS dust. The uncertainties involved in the derivation of both MODIS and MISR dust inevitably contribute to their differences too. However, it should be noted that although both methods have limitations, MISR dust discrimination is based on actual retrieved particle shape information, whereas the accuracy of MODIS dust is subject to the use of constant $f$ values and empirically calculated $f_{	ext{map}}$. Despite more physically robust separation of dust from in MISR, it has much less frequent sampling compared to MODIS, thus it is necessary to examine both data sets. Results based on these two complementary satellite data sets will provide more solid insight to the characteristics of the MJO-related dust and to some extent smoke variations.

4. MJO-Related Atlantic Dust and Smoke AOT Anomalies

4.1. MJO Analysis Methodology

For the MJO analysis and composite procedure, we use the multivariate empirical orthogonal function method introduced by Wheeler and Hendon (2004) and adopted widely by the MJO community (e.g., Waliser et al., 2009). In this method, eight distinct phases of the MJO cycle are determined by a pair of Real-time Multivariate MJO (RMM) index (RMM1 and RMM2) [Wheeler and Hendon, 2004]. Typically, when enhanced MJO convection is located over the equatorial Indian Ocean (phase 1 and 2) and western hemisphere (phase 8), persistent low-level westerly anomalies and enhanced precipitation are found over the equatorial Atlantic, whereas the reverse conditions are found when the MJO convection is located over the maritime continent (phase 4) and western Pacific (phase 5 and 6).

In this study, first, the intraseasonal anomalies of daily AOT are obtained by removing the climatological-mean seasonal cycle and filtering via a 30–90 day bandpass filter. Then, a composite MJO cycle is calculated by averaging the daily anomalies that occur within each phase of the MJO cycle. Only days having strong MJO activity ($\text{RMM1}^2+\text{RMM2}^2>1$) are considered for the MJO cycle composite.

4.2. MJO Composites of Dust and Smoke Anomalies

Figure 5 shows the number of strong MJO days ($\text{RMM1}^2+\text{RMM2}^2>1$) used for the eight-phase MJO cycle composite for both MODIS and MISR during 2002–2009 boreal winters. The total number of strong MJO events during this period is also shown in the upper right corner of each panel. For both instruments, the number of strong MJO events used for the composite is much less than the actual number of total events because of satellite retrieval sampling issues and the data rejection applied. Generally, the number of MODIS events ranges from about 10 to 45, about as three times more than MISR, due to the much wider MODIS swath.

The eight-phase MJO composite maps of MODIS total $\tau$, dust $\tau_{\text{dust}}$, and smoke $\tau_{\text{smoke}}$ anomalies are shown in Figures 6a, 7a, and 8a, respectively. The MISR counterparts will be
Figure 6. MJO composite maps of total $\tau$ anomalies (multiplied by 100) for (a) MODIS and (b) MISR over the tropical Atlantic Ocean. Nine-point spatial smoothing is applied.

Figure 7. Same as Figure 6, but for $\tau_{\text{dust}}$. 
discussed and compared with MODIS results later. Comparing Figures 6a and 7a, MODIS $\tau$ and $\tau_{dust}$ have very similar temporal evolution and spatial patterns, except that $\tau_{dust}$ anomalies are slightly smaller. The spatial correlation between $\tau$ and $\tau_{dust}$ anomalies for eight MJO phases is 0.89. For both $\tau$ and $\tau_{dust}$, strong negative anomalies (as large as about $-0.04$) are found over the equatorial Atlantic for MJO phases 1, 2, 3, and 8. In contrast, strong positive anomalies (up to 0.04) are found over the equatorial Atlantic for MJO phases 5–6. For MJO phase 4, strong positive anomalies occur to the north of the equator, whereas negative $\tau$ anomalies are found to the south. The converse is true for the MJO phase 7. The MJO composite maps of total $\tau$ greatly resemble those shown by Tian et al. [2011], with spatial correlation 0.52 (significant at the 99.9% level), and have very similar magnitudes. This resemblance suggests that the MJO-related total AOT anomaly patterns are robust and not sensitive to the data sampling (fewer but more carefully filtered samples are used in this study). The MJO-related $\tau_{smoke}$ anomalies are very weak, rarely exceeding 0.01 (Figure 8a). The spatial correlation between $\tau_{smoke}$ and $\tau$ anomalies is only about 0.21. The MJO composite maps of $\tau_{mar}$ anomalies are not shown because $\tau_{mar}$ is linearly dependent on the surface wind speed, thus the MJO-related $\tau_{mar}$ anomaly pattern in fact reflects the wind anomalies associated with the MJO, which were examined in Tian et al. [2011]. Furthermore, it is found that the magnitudes of the $\tau_{mar}$ anomalies are negligible: the strongest negative/positive anomalies are about $-0.004/0.004$ (figure not shown).

The eight-phase MJO composite maps of MISR total $\tau$, $\tau_{dust}$, and $\tau_{smoke}$ anomalies are shown in Figures 6b, 7b, and 8b. Again, it is found that the $\tau_{dust}$ anomalies are significantly larger than the $\tau_{smoke}$ anomalies. The $\tau_{dust}$ anomalies have very similar patterns to those of total $\tau$, with slightly smaller magnitude, whereas the $\tau_{smoke}$ anomalies are small and noisy. Further comparison between the MISR and MODIS $\tau$ and $\tau_{dust}$ results (Figures 6 and 7) indicates that overall they exhibit very similar temporal evolution and anomaly patterns despite the systematically larger MISR anomalies compared to those of MODIS, which is likely the outcome of the systematically larger MISR $\tau$ and $\tau_{dust}$ retrievals compared to MODIS over the tropical Atlantic Ocean, and possibly more robust dust identification for MISR, as discussed earlier. The rejection of all $\tau < 0.15$ data in MISR could also contribute to the above difference; however, MISR $\tau$ anomalies with no data rejection do not show evident differences (Figure not shown). Besides the magnitude difference, MISR $\tau$ and $\tau_{dust}$ anomalies are also noisier than the MODIS anomalies due to lower sampling. Furthermore, although MISR $\tau_{smoke}$ anomalies are small too, they are notably stronger than the MODIS $\tau_{smoke}$ anomalies. Nevertheless, the consistency between the MODIS and MISR results shown in Figures 6–8 demonstrate that dust is the dominant aerosol component on the intraseasonal time scale, and the MJO-related dust anomalies are robust, as seen from two independent sets of satellite observations having different strengths and limitations.

4.3. Sensitivity of MODIS Results

[39] The Kaufman et al. method for computing $\tau_{dust}$ and $\tau_{smoke}$ using MODIS $\tau$ and $f$ based on equations (3) and (4) is straightforward; however, large uncertainties in the computed $\tau_{dust}$ and $\tau_{smoke}$ are expected for at least several
Sensitivity tests and the control case, we average the composite aerosol anomalies in each MJO phase over a representative region (30°W–15°W, EQ–15°N) where the aerosol shows strong intraseasonal variations. First, the MJO cycles of the τ, τ_dust, τ_smoke, and τ_mar anomalies based on the control case (see Figures 6a, 7a, and 8a for τ, τ_dust, and τ_smoke anomalies) are shown in Figure 9a. It is found that the magnitude of τ_dust anomalies averaged over this region is as large as that of τ anomalies, whereas τ_smoke and τ_mar have very small magnitudes. Sensitivity test results for τ_dust and τ_smoke are shown in Figures 9b and 9c, respectively. For τ_dust, overall, the colored lines (nine sensitivity tests) cluster around the solid black curve (control case), and the MJO cycles of τ_dust anomalies based on different sensitivity tests show a coherent evolution. This suggests that the MJO associated τ_dust anomalies over the tropical Atlantic are quite robust despite the uncertainties in using constant f values. Nevertheless, the lines spread in some phases. The result based on Yu et al. [2009] is most different from the other cases, probably because the f values in the other studies follow the sequence f_mar < f_dust < f_smoke despite the deviations, whereas the order of f_mar and f_dust is reversed in the Yu et al. [2009] study, i.e., the size distribution for maritime aerosol is assumed by Yu et al. [2009] to contain a larger AOT fraction in the fine mode than that for dust. The spread in group 2 is naturally larger than that in group 1. Furthermore, the τ_dust anomalies are more sensitive to f_dust as indicated by the larger deviations of the dashed cyan and blue curves relative to the control case in phases 2, 3, and 6.

The spread of τ_smoke anomalies is quite large compared to their magnitude (Figure 9c). Thus, given the large uncertainties involved in the τ_smoke derivations and the small and noisy τ_smoke anomalies found in Figure 8, the numbers shown in Figure 9c should be more regarded as giving the signs of the MJO-related anomalies in each phase than the precise quantification of the anomalies. Nevertheless, the overall MJO cycle of τ_smoke anomalies is consistent among the different cases.

4.4. Modulation of Dust and Smoke Anomalies by the MJO

The MJO is characterized by eastward propagation of enhanced or suppressed zonal wind and precipitation anomalies. Over the tropical Atlantic, low-level wind and precipitation are two critical parameters that lead to aerosol variations through their influence on aerosol emission, transport, and deposition. Tian et al. [2011] examined the lag-correlation between the total τ anomalies and the low-troposphere zonal wind and precipitation anomalies related to the MJO. They found that the MJO modulates the tropical Atlantic aerosol primarily through the zonal wind, which can last as long as one MJO phase. The precipitation also exhibits some modulation on the aerosol, but the correlation is much weaker compared to the wind impact. In this section, we further examine the modulation of the MJO of dust and smoke aerosols through the lower-tropospheric zonal wind and precipitation.

Linear lag correlation between the MODIS τ_dust anomalies (Figure 7a) and the low-troposphere zonal wind and precipitation anomalies [Tian et al., 2011, Figures 8 and 6] for the composite MJO cycle over the tropical Atlantic region is
shown in the first two columns of Figure 10. It is found that the dust anomalies are strongly and negatively correlated with the wind anomalies over most of the tropical Atlantic region for lag 0, suggesting that the trade wind anomalies could modulate dust aerosol through westward transport, dust emission and/or dry deposition: trade winds produce a prevailing easterly component over the tropical Atlantic, thus a positive zonal wind anomaly means a relaxed trade wind, which corresponds to weakened westward dust transport, and possibly stronger dry deposition as well as weaker dust emission over the source region, all resulting in negative aerosol concentration, and vice versa. This strong negative correlation is also evident when the zonal wind anomalies lead the dust anomalies by one MJO phase, indicating that the wind modulation lasts as long as one MJO phase, which is consistent with what was found by Tian et al. [2011]. Negative correlations between dust and precipitation anomalies are also found in lag 0 and lag −1, but the correlations are much weaker and mostly confined within the convective equatorial region compared to the correlation with the wind anomalies, suggesting moderate MJO modulation on dust aerosol through wet scavenging. Note that some scattered positive correlations are found especially along the South American coast, which might be related to the cloud contamination to the aerosol retrieval.

In contrast with the dust case, examination between the smoke anomalies and the wind and precipitation anomalies suggests that the MJO exhibits very little modulation of the smoke aerosol (right two columns of Figure 10).

Figure 10 shows that dust aerosol is heavily influenced by the MJO through both dynamical and thermodynamical processes, whereas the smoke is not. This might explain the relative magnitude of dust and smoke aerosol anomalies on the intraseasonal time scale. As shown earlier, the MJO composite τ_{dust} anomalies are as large as the total τ anomalies (compare Figures 9a and 10a, and the black and red curves in Figure 9a). Although the MJO related τ_{smoke} anomalies are noisy and small, thus difficult to measure, they are about 10% of the total τ anomalies, as roughly inferred from Figure 9a. Because the above numbers are based on the composite of strong MJO events only, we further examine the standard deviations of the 30–90 day filtered dust and smoke anomalies, together with those of the total τ and τ_{mar} (Figure 11). It is found that the magnitude of τ_{dust} anomalies is comparable to or even larger than that of the total τ, whereas the τ_{smoke} anomaly represents about 15% compared to that of total τ, and τ_{mar} shows very little intraseasonal variability, which is overall consistent with what we found based on the composite of strong MJO events. Note that although the intraseasonal τ_{dust}, τ_{smoke}, and τ_{mar} anomalies add up to the intraseasonal total τ anomaly for each day, their variance does not sum to the variance of total τ for this case, suggesting the individual aerosol components are not
completely independent of each other on the intraseasonal time scale, so the nonlinear terms do not vanish.

Recall that $\tau_{\text{dust}}$ is about 40–70% of total $\tau$, and $\tau_{\text{smoke}}$ and $\tau_{\text{mar}}$ contribute another around 25% each, in terms of winter climatological mean (Figures 1c, 1e, and 1g), which is quite different from the relative magnitude of $\tau_{\text{dust}}, \tau_{\text{smoke}},$ and $\tau_{\text{mar}}$ on the intraseasonal time scale as shown above. Thus, we find that although the dominance of dust over smoke and marine aerosols on the intraseasonal time scale can be inferred from their magnitudes in the climatological mean, the former is not necessarily proportional to the latter. Note that the 25% V.S. 15% difference in smoke on the climatological and MJO-related scales does not lead to big difference in absolute AOT value, actually is estimated as small as about 0.0015. Such a small value is likely smaller than the uncertainty in an individual retrieval of the smoke AOT. Nevertheless, it should be noted that these percentages are obtained by averaging all data during 2002–2009, thus they should be accurate enough in terms of telling the right direction of the change from the climatological scale to the MJO-related scale: For smoke, it is a decrease from 25% to 15%, while it is an increase from about 40–70% to about 100% for dust. Compared to smoke, dust shows an even stronger fluctuation on the intraseasonal time, which is due to the fact that dust is more susceptible to the MJO, as it is more heavily influenced by both the trade wind and precipitation anomalies associated with the MJO.

5. Summary and Conclusions

Previous studies [Tian et al., 2011] found significant intraseasonal variability related to the MJO in the total

![Figure 11. Standard deviation of MODIS aerosol intraseasonal anomalies (30–90 day bandpass filtered) during 2002–2009 boreal winter for (a) total $\tau$, (b) $\tau_{\text{dust}}$, (d) $\tau_{\text{smoke}}$, and (f) $\tau_{\text{mar}}$, and their percentages to $\tau$ ((c) $\tau_{\text{dust}}$, (e) $\tau_{\text{smoke}}$, and (g) $\tau_{\text{mar}}$). Nine-point spatial smoothing is applied.](image-url)
column AOT over the tropical Atlantic region. Aerosol over the tropical Atlantic is primarily a mixture of mineral dust, biomass burning smoke, and marine aerosol. Given the different roles in the radiative forcing and cloud formation processes played by these three aerosol types, as well as the potential predictability of the MJO extending to 2–4 weeks, it is of great interest to further examine the MJO-related variability for individual aerosol types, especially dust and smoke aerosols. In this study, we used two independent and complementary satellite data sets, MODIS and MISR, to investigate the intraseasonal variability of individual aerosol components over the tropical Atlantic.

Daily MODIS/Aqua total AOT and fine-mode fraction measurements are used to derive daily \( \tau_{\text{dust}} \) and \( \tau_{\text{smoke}} \) following the method of Kaufman et al. [2005a, 2005b]. This method contains considerable uncertainties, due to the assumption of constant \( f_{\text{dust}}, f_{\text{smoke}}, f_{\text{mar}} \) and the empirical calculation of \( \tau_{\text{mar}} \). Strict data rejection has been applied to use Kaufman’s formula in a safer way. With MISR’s sensitivity to aerosol particle shape, dust and smoke aerosols can be distinguished using the MISR aerosol nonspherical fraction. Because MISR sensitivity to shape diminishes when aerosol concentration is low, only \( \tau > 0.15 \) data are used to compute \( \tau_{\text{dust}} \) and \( \tau_{\text{smoke}} \) for MISR. The examination of both data sets is necessary, and results from MODIS and MISR are complementary: MODIS provides better spatial coverage, and MISR dust is derived from actual aerosol property retrieval rather than assumed aerosol-type-specific factors. MODIS and MISR show a very similar dust and smoke winter climatologies. \( \tau_{\text{dust}} \) is found to be the dominant aerosol component over the tropical Atlantic. It is largest over the eastern equatorial Atlantic (about 0.35) and gradually decreases toward the west. The contribution of \( \tau_{\text{dust}} \) to total \( \tau \) ranges from 40% to 70%, considerably larger than \( \tau_{\text{smoke}} \), which contributes approximately 25%, and \( \tau_{\text{mar}} \) contributes less than 20%. The daily MODIS and MISR \( \tau_{\text{dust}} \) distributions are highly correlated overall, with the correlation coefficients typically about 0.7 over the North Atlantic, but much smaller or even negative over the South Atlantic. MISR \( \tau_{\text{dust}} \) is found to be systematically greater than the coincident MODIS \( \tau_{\text{dust}} \), and this difference can be traced to the AOT difference between them. The consistency of the MODIS and MISR dust and smoke aerosol climatologies and daily variations lend confidence to our use of these data sets to investigate the relative contributions of dust and smoke aerosols to the total AOT variation associated with the MJO. However, the identification of smoke is much less certain than that of dust, because discrimination among fine-mode sea salt, sulfate, and smoke particles depends on assumptions for both MODIS and MISR, whereas the MISR dust discrimination is based on retrieved particle shape.

For MODIS, the MJO composite maps of \( \tau_{\text{dust}} \) anomalies are very similar to those of \( \tau \) anomalies, and are of comparable magnitude. Furthermore, the variance of the \( \tau_{\text{dust}} \) anomalies on the full intraseasonal time scale is found to be comparable or even bigger than that of the \( \tau \) anomalies. In contrast, the MJO-related \( \tau_{\text{smoke}} \) anomalies are rather small, barely exceeding 0.01, and the \( \tau_{\text{mar}} \) anomalies are negligible. The sensitivity study further shows that the MJO-related \( \tau_{\text{dust}} \) and \( \tau_{\text{smoke}} \) anomalies are quite robust, even when we perturb the \( f \) values by their uncertainty ranges or use different sets of \( f \) values from several independent studies.

Similarly, MISR also shows that the MJO composite maps of \( \tau_{\text{dust}} \) anomalies are very similar to those of the \( \tau \) anomalies, whereas the MJO-related \( \tau_{\text{smoke}} \) anomalies are rather small. The composite MJO cycle of \( \tau_{\text{dust}} \) anomalies from MISR over the tropical Atlantic Ocean is consistent with the MODIS results, although the anomalies are much noisier due to less frequent sampling. The magnitude of MISR anomalies is again found to be systematically larger than that of MODIS. The MJO-related \( \tau_{\text{smoke}} \) anomalies in MISR are overall slightly larger than for MODIS, but still much smaller compared to the \( \tau_{\text{dust}} \) anomalies.

The consistency between the MODIS and MISR \( \tau_{\text{dust}} \) and \( \tau_{\text{smoke}} \) anomalies in terms of the evolution of the MJO cycle and the spatial pattern of anomalies suggests that dust aerosol is the dominant component on the intraseasonal time scale over the tropical Atlantic Ocean, and supports the other main conclusions of this study.

Although dust contributes about 40–70% to total \( \tau \) in terms of climatological mean, its intraseasonal variability is comparable to or even larger than the intraseasonal variability of total \( \tau \). The smoke intraseasonal variability is about 15% of that for total \( \tau \), even though its climatological mean contributes to about 25% of the total \( \tau \). This suggests that the MJO does not influence dust and smoke in the exactly same way. Examination of the lag correlation between dust and smoke aerosol anomalies with the low-level zonal wind and precipitation anomalies support this speculation: dust is more susceptible to the MJO as is more heavily influenced by the trade wind and precipitation anomaly associated with the MJO. The modulation of the MJO-related wind anomalies on the dust anomalies can last as long as one MJO phase.

The observational results obtained from two complementary satellite data sets can be used to evaluate chemical transport models and help in model development. Furthermore, the findings of this work have broader implications, related to the predictability issue. As we know, predictability beyond the synoptic time scale relies merely on our knowledge of the semi-periodic phenomena such as the MJO. Given that potential prediction of the MJO is extended to a few weeks currently, our finding suggests that dust has a potential prediction time scale up to a few weeks, too. This adds potential guidance to the prediction of phenomena affected by dust, such as dust storms, Atlantic tropical cyclogenesis, Atlantic tropical cyclone evolution, ocean fertilization, and so on.

Appendix A: MODIS Data Rejection

In this work, data rejection is performed for the MODIS \( \tau \) and \( f \) measurements for the reasons described below.

The dependence of \( \tau_{\text{dust}} \) and \( \tau_{\text{smoke}} \) on \( \tau \) and \( f \) based on equations (3) and (4) is illustrated in Figure A1a and A1b, where \( \tau_{\text{mar}} \) is set to 0.06 (the approximate global mean value of \( \tau_{\text{mar}} \)). It is seen that the larger the \( \tau \), the larger \( \tau_{\text{dust}} \) and \( \tau_{\text{smoke}} \). Also, as \( f \) increases, \( \tau_{\text{smoke}} \) increases and \( \tau_{\text{dust}} \) decreases. It is shown that equations (3) and (4) can produce reasonable (nonnegative) values for both \( \tau_{\text{dust}} \) and \( \tau_{\text{smoke}} \) only when the paired \( \tau \) and \( f \) measurements fall within a limited region bounded by the white dashed lines \( \tau_{\text{dust}} = -0.03 \)
t and smoke = 0.03). We relax the limit of t dust or t smoke to 0.03 instead of 0 because the uncertainties of the MODIS AOT are 0.03 over ocean, thus a small negative AOT up to 0.03 is regarded as indistinguishable from the value 0 \cite{Remer et al., 2005}. Note that calculated t dust and t smoke in the range of 0 to 0.03 have been set to zero. Outside this region, either t dust or t smoke is too negative, and correspondingly, t smoke or t dust would be larger than the total AOT, which is not physical. Therefore, the \(\tau\) and f measurements giving rise to such t dust and t smoke values are rejected. In addition, \(\tau\) measurements greater than 2 are excluded, given the large uncertainties from possible cloud contamination \cite{Zhang et al., 2005}.

\[ [57] \] The count distribution of paired MODIS \(\tau\) and f measurements as a function of total \(\tau\) and f according to equations (3) and (4). Here \(f_{mar} = 0.3, f_{dust} = 0.5, f_{smoke} = 0.9, \) and \(\tau_{mar} = 0.06, \) as indicated by four white straight lines. The white dashed lines indicate \(\tau_{dust} = -0.03\) and \(\tau_{smoke} = -0.03.\) Contour interval: 0.1. (c) The count of paired MODIS \(\tau\) and f measurements with respect to \(\tau\) and f over the tropical Atlantic (20\(^\circ\)S–30\(^\circ\)N, 60\(^\circ\)W–20\(^\circ\)E) for 2002–2009 boreal winters. Superimposed solid black lines are the two white dashed lines in Figures A1a and A1b corresponding to \(\tau_{dust} = -0.03\) and \(\tau_{smoke} = -0.03.\)

\[ [58] \] Before data rejection, MODIS aerosol observations are available for about 70% of the days during 2002–2009.

\[ [59] \] Before data rejection, MODIS aerosol observations are available for about 70% of the days during 2002–2009.
boreal winters over the clear subtropical North Atlantic, and decreases to about 35% over the cloudy equatorial Atlantic. The data rejection procedure described above rejects somewhat less than half of the observations. The most frequent rejections occur over the subtropical oceans, possibly due to the ambiguity between sea salt and dust aerosol over those regions.


